



Two-sided asymmetric subduction; implications for tectonomagmatic and metallogenic evolution of the Lut Block, eastern Iran

R. Arjmandzadeh^{1*}, M.H. Karimpour¹, S.A. Mazaheri¹, J.F. Santos², J.M. Medina², S.M. Homam¹

1 -Department of Geology, Ferdowsi University of Mashhad, Iran

2 -Department of Geosciences, GeoBioTec Research Unit, University of Aveiro, Portugal

Received: 25/9/2010, in revised form: 15/1/2011

Abstract

West directed subduction zones show common characteristics, such as low structural elevation, deep trench, steep slab and a conjugate back-arc basin that are opposite to those of the east directed subduction zones. The tectonomagmatic and metallogenic setting of the Lut Block is still a matter of debate and several hypotheses have been put forward. Despite some authors denying the influence of the operation of Benioff planes, the majority propose that it occurred beneath the Afghan Block, while others consider that oceanic lithosphere was dragged under the Lut Block. Cu-Au porphyry deposits seem to occur in an island arc geotectonic setting during the middle Eocene while Mo-bearing deposits are coincident with the crustal thickening during Oligocene. We introduce new trace element and isotope geochemical data for granitoids and structural evidences testifying the two-sided asymmetric subduction beneath both Afghan and Lut Blocks, with different rates of consumption of oceanic lithosphere.

Keywords *asymmetric subduction, tectono-magmatic, Lut Block, Afghan Block, isotope geochemical data.*

Introduction

The Lut Block has been considered one of the nine structural zones of Iran since the work of Stöcklin [1]. This block is bounded to the east by the Nehbandan and associated faults, to the north by the Doruneh and related faults (Sabzevar Zone), and to the west by the Nayband Fault (Fig. 1). The South Jazmourian fault, in the south of Sahand-Bazman magmatic arc, probably marks the southern limit of the block [2]. Some authors denied the influence of a subduction zone and attributed the mineralization in the Lut Block to an extensional geotectonic zone [3, 4, 5, and 6]. However, Saccani et al. [7] studied the ophiolitic complex of Eastern Iran, between the Lut and the Afghan continental blocks, and considered that the subduction of oceanic lithosphere played a major role and that it should have taken place beneath the Afghan Block. On the other hand, Eftekharnajad [8] proposed that magmatism in the northern Lut area resulted from subduction beneath the Lut Block. Additionally, Berberian [9] showed

that igneous rocks of this block have calcalkaline signatures. The accretionary prism-forearc basin polarity, the structural vergence and younging of the accretionary prism to the southwest are consistent with a northeast-dipping subduction [10]. The purpose of these studies is to present major, trace element and Sr-Nd isotope geochemical data for many Cu-Au-Mo porphyry bearing granitoids (Fig. 1) and some structural evidences to make a discussion on the origin of magma and the tectonomagmatic evolution of the Eastern Iran zone. Recently, asymmetric subduction models have been discussed for situations similar to that of the Lut Block [11 and 12]. This type of hypothesis will be discussed in the present work, taking into account that subduction related magmatism occurs in both Lut and Afghan Blocks but also that the structural evidence alone would point to a single subduction under the Afghan Block.

¹ corresponding author: arjmand176@gmail.com

Archive of SID

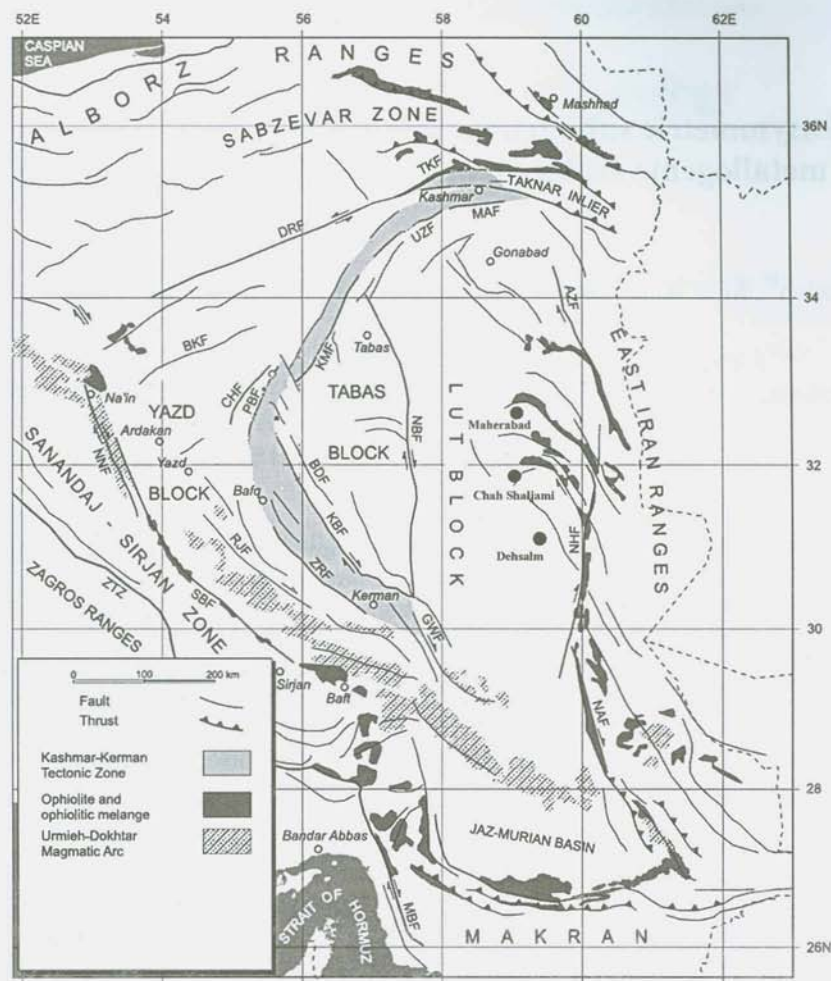


Fig. 1. The structural map of Central-East Iran and its constituent crustal blocks [13, 14, 15, 16 and 17]. AZF: Abiz Fault; BDF: Behabad Fault, BKF: Biabanak Fault, CHF: Chapedony Fault, DRF: Doruneh Fault, GWF: Gowk Fault, KBF: Kuhbanan Fault, KMF: Kalmard Fault, MAF: Mehdiabad Fault, MBF: Minab Fault, NAF: Nostratabad Fault, NHF: Nehbandan Fault, NNF: Na'in Fault, RJF: Rafsanjan Fault, SBF: Shahre-Babak Fault, TKF: Taknar Fault, UZF: Uzbak-Kuh Fault, ZRF: Zarand Fault, ZTZ: Zagros Thrust Zone.

Materials and methods

- 110 thin and 40 polished sections from various geological units of Dehsalm and Chah-Shaljami prepared for petrography and mineralography studies.
- 1091 soil, 187 rock and 41 mineralized samples from Dehsalm area analyzed for 45 elements at the AMDEL laboratory.
- 43 stream sediments, 7 heavy minerals, 26 mineralized samples from Chah-Shaljami area analyzed at the Exploitation Research Center of Iran, Karaj.

- 26 mineralized samples from Chah-Shaljami area analyzed for 28 elements at the ALS-Chemex laboratory.
- 25 samples of the intrusive rocks were analyzed for major elements by wavelength-dispersive X-ray fluorescence (XRF) spectrometry using fused discs and the Philips PW 1410 XRF spectrometer at the Ferdowsi University of Mashhad, Iran.
- 20 of these samples were analyzed for trace elements using ICP-MS, following a lithium metaborate/tetraborate fusion and nitric acid total

Archive of SID

digestion, in the Acme Laboratories, Vancouver, Canada.

- Sr and Nd isotopic compositions were determined for 14 whole rocks and 4 mineral samples of plagioclase and biotite of the Dehsalm Chah-Shaljami granitoids at the Laboratório de Geologia Isotópica de Universidade de Aveiro, Portugal.

- 10 double polished samples are currently preparing for fluid inclusion studies at the Ferdowsi University of Mashhad, Iran.

Two-sided asymmetric subduction

Subduction zones can be analyzed in terms of a wide range of parameters, such as convergence rate, topographic and structural elevation of the related orogen, subsidence rate in the trench or fore deep, erosion rate, metamorphic evolution, magmatism, dip of the foreland monocline, depth and geometry of the decollement planes that generate the accretionary prism and the belt of the upper plate, the thickness and composition of the upper and lower plates, gravity, magnetic and heat flow anomalies, seismicity and slab dip. Therefore, there is a long list of parameters, which are relevant to the geometry and evolution of each particular subduction zone. Generally, the west directed subduction zones like Barbados, Apennines, Marianas and Tonga are characterized by low topography and low structural elevation, a deep trench or fore deep with high subsidence rates, generally a steep slab, an accretionary prism mostly composed by the shallow rocks of the lower plate and a conjugate back arc basin. In contrast, the east directed subduction zones (e.g., Andes) or north-east (Himalayas, Zagros) exhibit opposite signatures such as high structural and morphological elevation,

generally no back arc basin, shallower trench or fore deep with lower subsidence rate, deeply rooted thrust planes affecting the whole crust and lithospheric mantle, ultra-high pressure rocks and wide outcrops of metamorphic rocks, and dominantly shallower dip of the slab. Such an asymmetry in the slab dip has been demonstrated in Fig. 2.

The W-directed subduction zones are generally faster because they have the subduction hinge generally moving away with respect to the upper plate, and converging relative to the lower plate faster than the upper plate. Therefore, these subduction zones should supply much larger volumes to mantle recycling than the opposite subduction zones (Fig. 3).

West-directed subduction should be hotter because it involves a thicker section of asthenosphere that is generally assumed $>1300^{\circ}\text{C}$. This can explain why it has lower P and S seismic velocity with respect to the E- or NNE directed subduction zones (Fig. 4).

Tatsumi and Eggins [22] have shown a correlation between convergence rate and volumes of magmatism along subduction zones. The larger volumes of subduction predicted along W-directed slabs should favour the formation of a greater amount of arc-related magma, and generation of large volumes of oceanic crust in the backarc setting for a number of reasons: 1) the larger subduction rate should generate also more abundant slab dehydration, lowering the melting temperature; 2) the thicker and hotter (asthenospheric) mantle wedge should have a thicker column of potential melting; 3) faster slab entering means also larger shear heating.

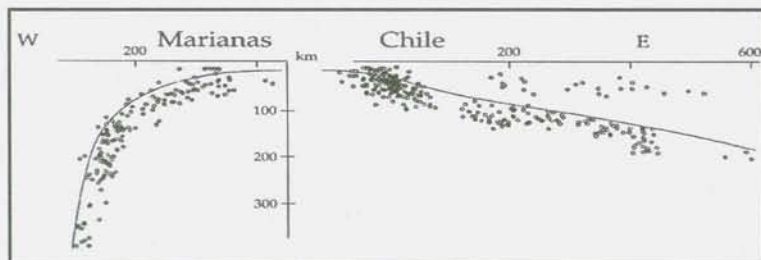


Fig. 2. Ipocenters of the Marianas and Chile subduction zones in the Pacific [18]

Archive of SID

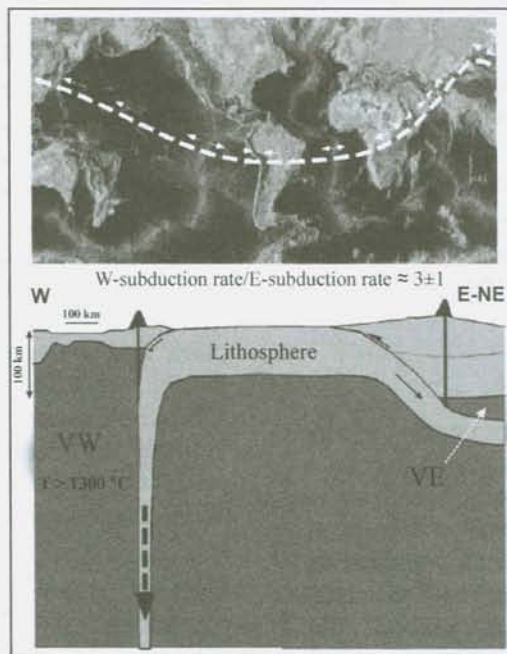


Fig. 3. The volumes recycled along W-directed subduction zones are about 2–3 times higher than along the opposite settings due to the aforementioned kinematic constraints. Moreover, the asthenospheric wedge above slabs is much thicker along W-directed subduction zones (VW) with respect to the E–NE directed subductions, if any (VE) modified after [19].

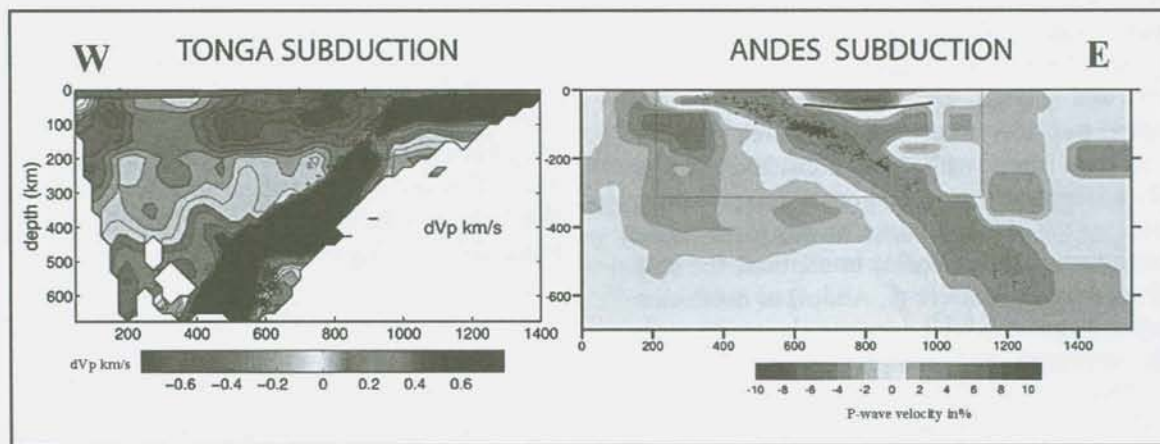


Fig. 4. Comparison between the Vp tomography of the Tonga and Andean subduction zones. Note the much slower velocities in the mantle wedge of the Tonga subduction [20] with respect to the opposite Andean setting [21].

Petrogenesis

Dehsalm and Chahjami igneous rocks are classified as high-K calc-alkaline to shoshonite rocks on Peccerillo and Taylor (Fig. 5) [23]. Linear increase of K_2O content from high K calc-alkaline to

shoshonitic series is related to the crystal fractionation as a result of crystallization of plagioclase, amphibole and pyroxene. The magnesian numbers ($\#Mg = 100 \times MgO / [MgO + Fe_2O_3]$, molar) are moderately high, ranging from 41.5 up to 55.5 and 40.3 up to

Archive of SID

60.9 respectively for Dehsalm and Chah-shaljami igneous rocks.

They are classified as volcanic arc granites on Pearce et al. (Fig. 6) [24]. Primitive mantle normalized trace element spider diagram [25] display strong enrichment in Large Ion Lithophile Elements (LILE) such as Rb, Sr, Ba, Zr and Cs; and depletion in some High Field Strength Elements (HFSE) e.g. Nb, P and Y (Fig. 7.a, b). The enrichment of the LILE and depletion of the HFSE and Heavy Rare Earth Elements (HREE) are the basic characteristics typical of the magmatism in a subduction belt [26] and calc-alkaline volcanic arcs. Their high Sr and low Nb, Ta and Ti contents are thought to be due to the absence of plagioclase and presence of Fe-Ti oxides in the residue [27].

All rocks show negative Nb anomalies as well as several other negative anomalies (P, Ta and Ti) that are also characteristic of subduction-related magmatism [28]. The negative peak of Nb is representative of arc zone and its depletion intensity might be related to the crustal effect [29]. The low

Nb and Zr contents of the rocks argue more for a calc-alkaline affinity; and low contents of Y, Nb, Ta and Yb indicate volcanic arc granites. Furthermore, low molar ratios $Al_2O_3/(Na_2O+K_2O+CaO) < 1.1$ and low Rb/Sr ratios with the mean of 0.15 and 0.19 for Dehsalm and Chah-Shaljami intrusives indicate that these rocks are I-type granitoids.

Strontium/Y and La/Yb ratios of Dehsalm and Chah-Shaljami intrusives are 31.6-72.2, 21.5-33.5; 19.7-67 and 21.4-33.7, respectively covering geochemical characteristics of adakites and are evident on Sr/Y-Y and La/Yb-Yb discrimination diagrams [30 and 31]. In the Sr/Y-Y diagrams (Fig. 8a), many samples are located in both adakitic and normal arc volcanic rocks fields and many out of the both fields. Samples out of the both fields have Sr/Y content similar to that of adakites but Y slight enrichment by minor crustal contamination and/or crystal fractionation lead to such a position. Such samples show more adakitic affinities on La/Yb diagram (Fig. 8b).

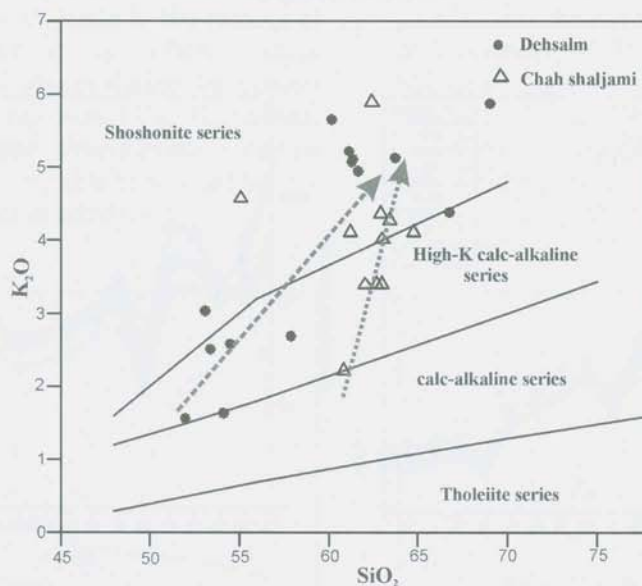


Fig 5. K_2O vs. SiO_2 plots for the different igneous suites in the Dehsalm and Chah-Shalghami. Peccerillo and Taylor [23].

Archive of SID

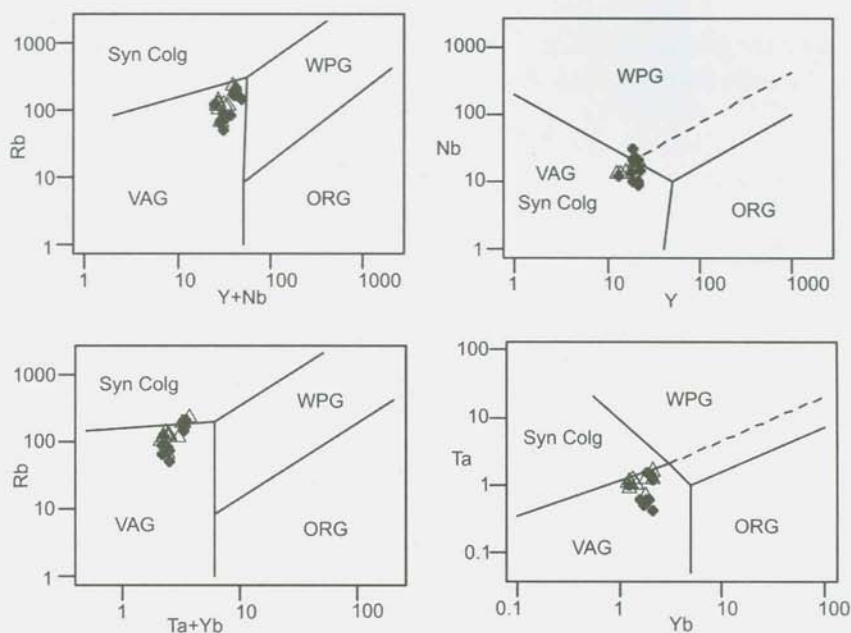


Fig. 6. Geotectonic diagrams of Pearce et al. [24]; for the different igneous suites of Dehsalm and Chah-Shaljami. The symbols like Fig. 5.

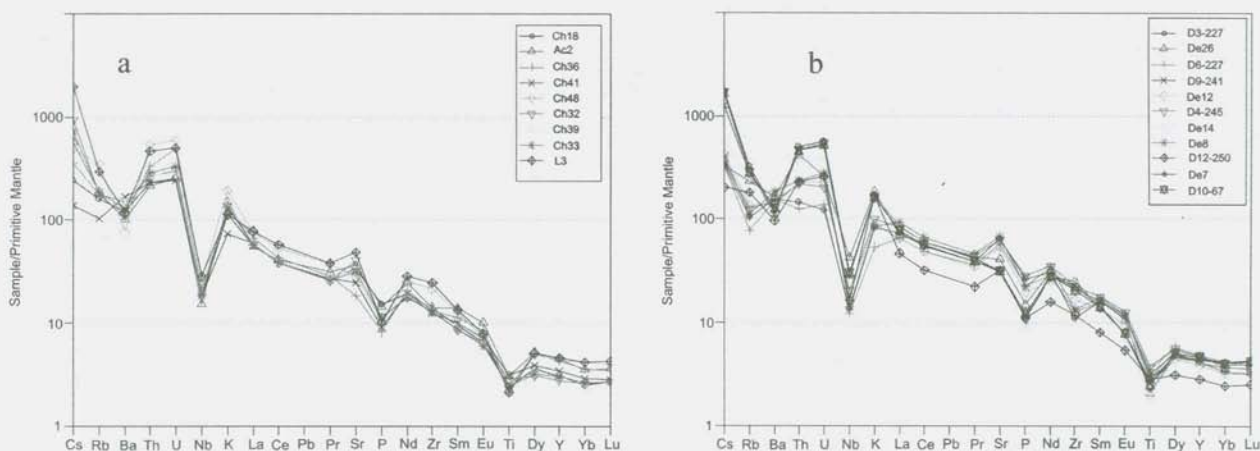


Fig. 7.a, b. Primitive mantle normalized trace element spider diagram [25] for the Chah-Shaljami and Dehsalm intrusives.

Archive of SID

Origin of the adakitic magmas

The main geochemical features of Dehsalm and Chah-shaljami intrusives like high Sr/Y and La/Yb ratios, low abundance of HFSE and HREE are similar to adakites. These characteristics are satisfied by melting of garnet amphibolite or eclogite facies rocks, which may be found at the base of thickened (>40 km) continental crust or in eclogitized subducting oceanic crust. On the other hand, garnet is also a stable phase in subcontinental mantle lithosphere as well as asthenosphere [33]. In addition to slab melts, there are different origins such as assimilation-fractional crystallization (AFC) processes [34], melting of mantle peridotite under hydrous conditions [35], partial melting of thickened lower crust [36 and 37] or delaminated mafic lower crust [37, 38, 39 and 40] determined for adakitic compositional features. The low contents of Y and Yb and high ratios of Sr/Y and La/Yb can be attributed to the retention of Y and HREE in residual garnet, or fractionation of garnet and hornblende. High Sr content is attributed to absence of significant plagioclase fractionation. The low HREE content of adakites is classically interpreted as reflecting the presence of garnet and hornblende in the residue of partial melting of their source, whereas these minerals are not residual phases during the genesis of typical calc-alkaline magmas [41]. Hornblende and/or Fe-Ti oxides (rutile, ilmenite) are common residual minerals, thus, being able to account for Ti-Nb-Ta negative anomalies in adakites.

All measured isotopic ratios have been corrected for an age of 33 Ma based on the Rb-Sr mineral-whole rock internal isochron of samples CH33 and D3-227. The plot of the Dehsalm and Chah-Shaljami samples on the $\epsilon_{\text{Nd}}-(^{87}\text{Sr}/^{86}\text{Sr})_i$ diagram (Fig. 9a, b) shows that their compositions do not fit into an origin of the parental magmas by melting of thick lower crust or Cenozoic subducted oceanic crust as proposed for typical adakites. In contrast, they have Sr and Nd isotope compositions very similar to those of normal island arc basalts, pointing to melting in a mantle wedge followed by magmatic differentiation.

Adakitic magmas, whether derived directly from partial melting of the subducted oceanic slab (MORB) or from lower crustal mafic rocks, usually show characteristics of low Mg#<40 and high Na₂O (>4.3 wt%) rather than high K₂O, regardless of melting degrees [42]. Moderately high Mg# of Chah-Shaljami intrusives from 40.3 up to 60.9 can also be attributed to the mantle derived melts.

If the source is inferred to have contained rutile residue, the adakitic rocks would be characterized by elevated super-chondritic Nb/Ta ratios (chondritic ratio: 19.9 ± 0.6) and Zr/Hf ratios (chondritic ratio: 34.3 ± 0.3) with strongly decreasing Nb and Ta concentrations [43]. The Dehsalm and Chah-Shaljami adakitic rocks are characterized by relatively lower Nb/Ta ratios and Zr/Hf ratios compared to chondrite. The low Nb/Ta ratios indicate the presence of a low-Mg amphibole in the restite, such that Nb is retained in the restite compared to Ta [44].

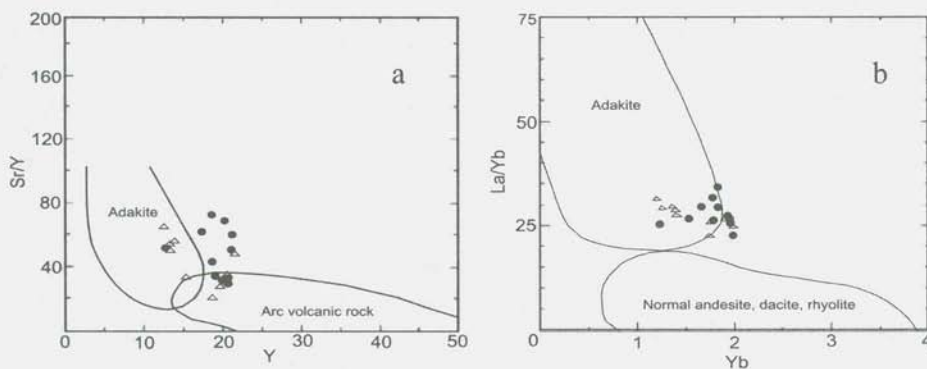


Fig 8. a,b: Y vs. Sr/Y and c, d: Yb vs. La/Yb diagram [32] used to differentiate adakitic magmas from typical calc-alkaline magmas for Chah-Shaljami intrusives. The symbols like Fig. 5.

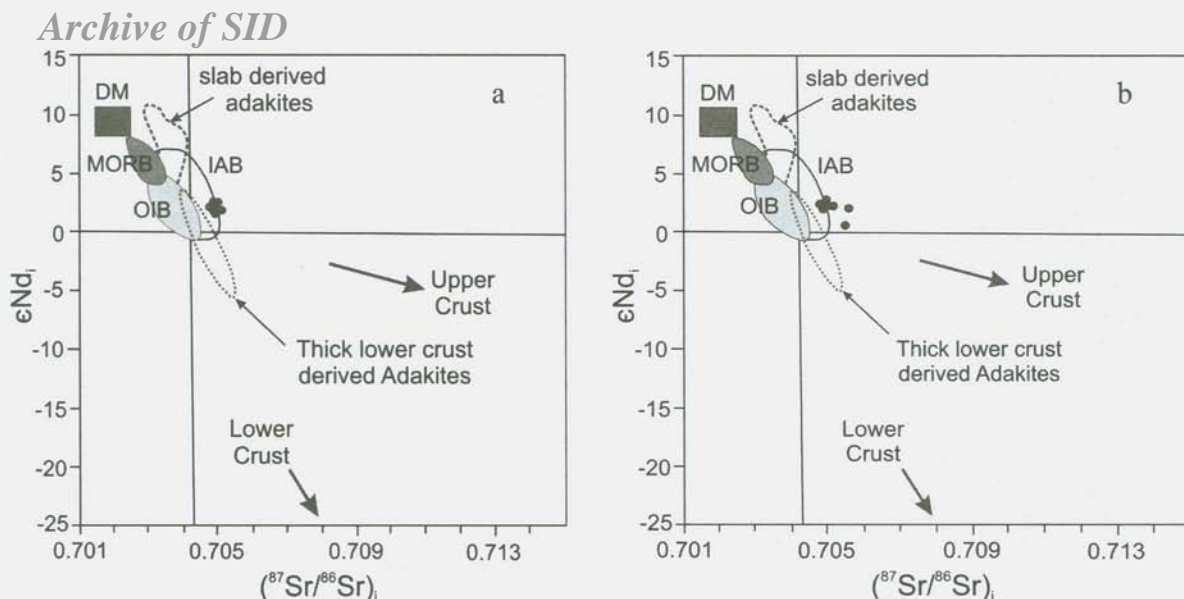


Fig. 9.a, b. ϵNd_t - $(^{87}\text{Sr}/^{86}\text{Sr})_t$ diagram for the Dehsalm and Chah-Shaljami intrusive rocks, respectively.

These results also show that rutile-bearing eclogites cannot serve as the source as partial melting of such eclogites produces melts with high Nb/Ta ratios [44]. Depletion in HREE and low Nb/Ta ratios requires both garnet and low-Mg amphibole in the restite.

Ionov and Hofmann [45] have shown from mantle xenoliths that amphiboles can have high K and very low Rb concentration while coexisting phlogopite is rich in both K and Rb. Thus, selective melting of phlogopite would account for high Rb content and correlated high K/Rb as for the Dehsalm and Chah-Shaljami intrusives. Metasomatism of mantle peridotite by slab-melts produces orthopyroxene, clinopyroxene, garnet, phlogopite, and richterite or pargasite [46]. Therefore, we suggest a high K, low Mg amphibole, phlogopite and garnet as the major components in the restite to rationalize geochemistry of the Dehsalm and Chah-Shaljami intrusives.

Tectonomagmatic and metallogenic implications

In Iran, Adakitic magmatism for the first time was introduced for Kharvana porphyry stock (NW Iran) by Arjmandzadeh and Alirezaei [47] and has recently been introduced within the Lut Block with the shoshonitic affinities [12].

Aftabi et al. [48] believe that the potassic magmatism in Iran might be related to porphyry

hydrothermal Cu-Mo-Au-Ag mineralization, which is a considerable exploration target in the area and merits further investigation. Various mineralizations like Cu-Mo-Au porphyry epithermal type deposits, Cu-Au-Ag IOCG type deposits, Cu and Au-Sb-Pb-Zn vein type deposits, Cu-Au massive sulfide type deposits, Sn-W-Au associated with reduced granitoids and Sn magmatic-skarn deposits formed during Jurassic to Tertiary magmatism phases in the Lut Block (Fig. 10). Sorkhkooh Cu-Mo and Gazu Cu porphyry type deposits are related to the subduction of oceanic lithosphere beneath the Lut Block [49]. Recent exploration and petrologic studies on the Lut Block volcanic-plutonic belt show capabilities and suitable targets for Cu-Au-Mo porphyry epithermal mineralization which are related to a subduction geotectonic setting [12 and 50].

While some authors deny the influence of the operation of Benioff planes [4, 5 and 6], most of the recent works consider that subduction of oceanic lithosphere had a major role in the tectonic evolution of this area. Until now, two types of geodynamic models considering subduction have been presented: some authors consider mostly structural evidences and propose that it occurred beneath the Afghan Block as shown in Fig. 11, [10, 51 and 7] whilst others consider the larger volumes of calc-alkaline

Archive of SID

magmatism and propose that oceanic lithosphere was dragged under the Lut Block (Fig. 12) [8, 9 and 49]. New trace element and isotope geochemical data, obtained for the Oligocene intrusives from Dehsalm and Chah-Shaljami areas, indicate that a subduction zone existed below the Lut Block [12]. On the other hand, structural evidence shows that subduction occurred beneath the Afghan Block. In order to explain the most important tectonomagmatic and metallogenic characteristics of the Lut Block, west-east directed subduction is discussed here.

As discussed above the E or NE directed slabs have slower sinking velocity than the opposite W-directed subduction zones (Fig. 3). West-verging subduction would be completely consumed before E or NE-verging subduction and consequently lead to the formation of structures that show evidences of one-sided subduction as reported between the Lut and Afghan Blocks (Fig. 13).

For example the accretionary prism-forearc basin polarity, the structural vergence and younging of the accretionary prism to the southwest are consistent with an NE dipping subduction beneath Afghan Block [10]. Therefore, structural evidences for E or NE-verging subduction does not preclude W-directed subduction under the Lut Block and a model with two-sided subduction has the advantage of also

accounting for the very important Tertiary calc-alkaline magmatism and associated mineralization within this block. Camp and Griffiths [52] reported the subduction of Sistan Ocean below the Afghan continental margin, as testified by the occurrence of both intrusive and extrusive Maastrichtian-Paleocene calc-alkaline rocks. The larger volumes of subduction predicted along W-directed slabs should favour the formation of greater amounts of arc-related magmas, as reported within the Lut Block, where voluminous Tertiary igneous rocks occur (Fig. 13). The ore-bearing intrusive rocks within the Maherabad Cu-Au porphyry deposit have been dated at 39 Ma (Middle Eocene) using zircon U-Pb and an island arc geotectonic setting proposed for the area [54]. The island arc geotectonic setting is coincident to the W-directed subduction (as shown in the Fig. 13) and the occurrence of Cu-Au mineralization during Eocene. According to Sengör and Natalin [55] the Ocean was closed in eastern Iran, between the Helmand and Lut plates in Oligocene–Middle Miocene. During this period, the crustal thickening occurred and a suitable geotectonic setting would result in the formation of Mo-bearing porphyry deposits as in cases of Oligocene Dehsalm and Chah-Shaljami deposits that have been dated at 33 Ma.

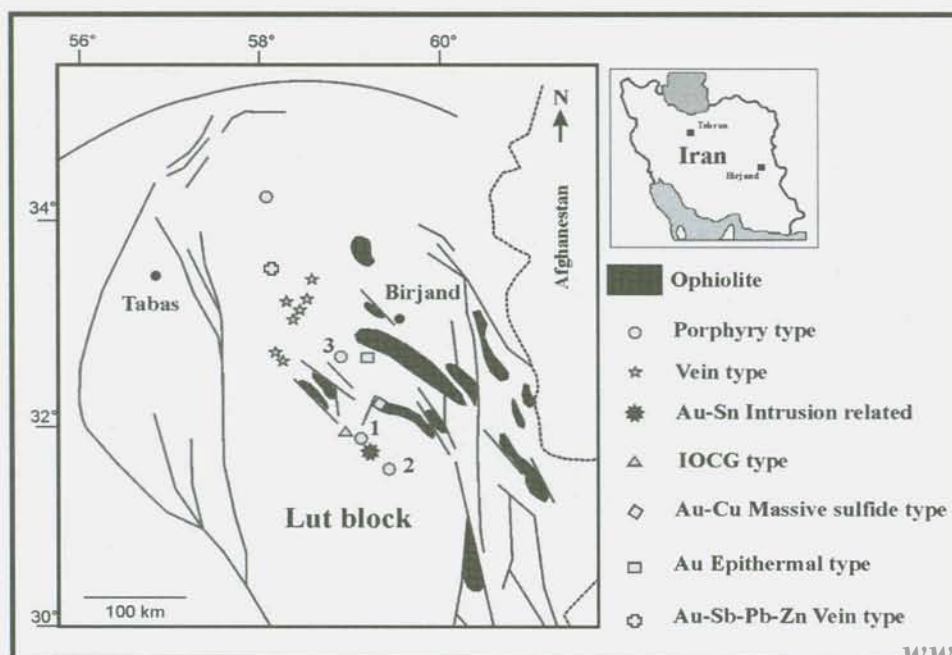


Fig. 10. Different types of mineralization occurrences within the Lut Block during Tertiary.

Archive of SID

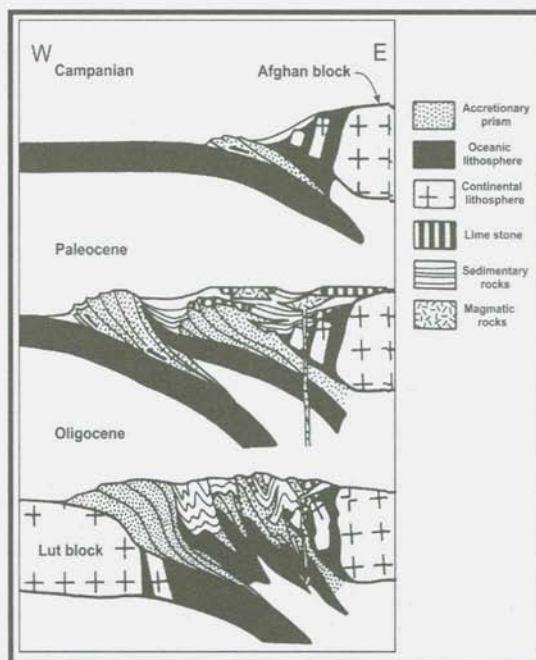


Fig. 11. Modified east verging subduction model of Eastern Iran [52, 10, 51, 53 and 7]. Such a geodynamic model cannot serve tectonomagmatic and metallogenic problems of the Eastern Iran.

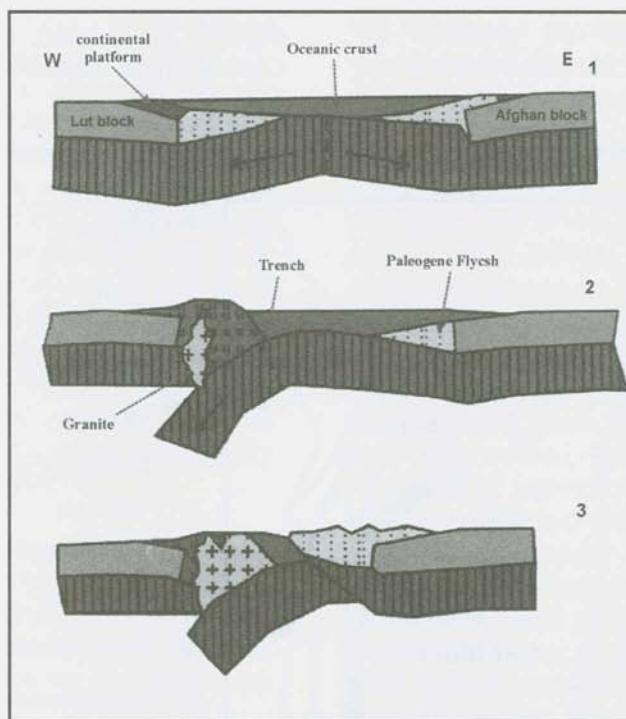


Fig. 12. Modified west verging subduction model of Eastern Iran [8]. Such a geodynamic model cannot serve tectonomagmatic and metallogenic problems of the Eastern Iran.

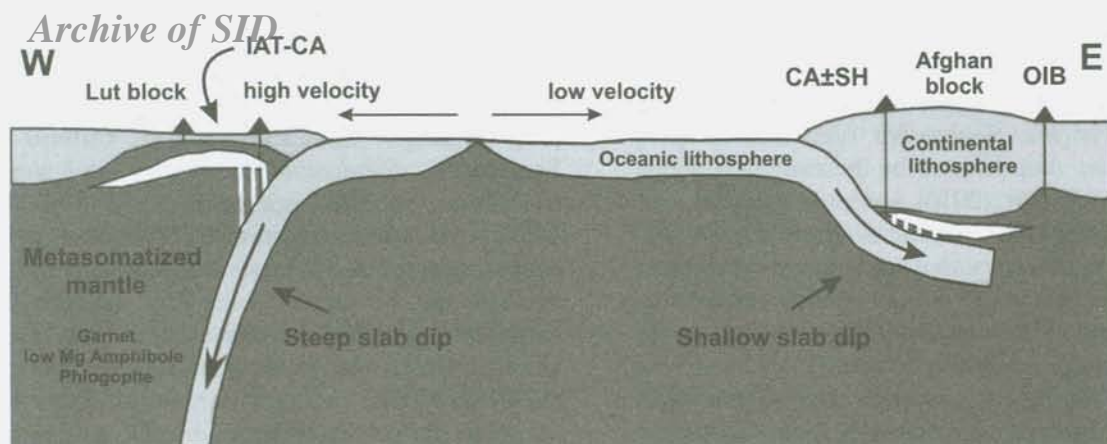


Fig. 13. Two-sided asymmetric subduction. This model is presented as a new hypothesis for the tectonomagmatic and metallogenic setting of the Lut Block. IAT: island-arc tholeiites; CA, SHO: calc-alkaline and shoshonitic series; OIB: basalts with ocean island or intraplate affinity.

Conclusions

Trace element features of Dehsalm and Chah-Shaljami granitoids show a typical magmatism related to a subduction zone, such as LILE-enrichment and marked Nb, Ta and Ti negative anomalies. The adakitic affinity of the Dehsalm and Chah-Shaljami rocks can be attributed to the presence of residual garnet and low-Mg amphibole in a mantle source. Isotope geochemistry and relatively high Mg# show that the parental magmas probably derived from partial melting of metasomatized peridotite in a supra-subduction mantle wedge. Copper-Au porphyry deposits seem to occur in an island arc geotectonic setting during Middle Eocene whilst Mo-bearing deposits are coincident with the crustal thickening during Oligocene. Such a geochemical data testify the West direction subduction beneath the Lut Block; along with the present structural evidences related to the east direction subduction beneath the Afghan Block we can consider a two-sided subduction. Two-sided asymmetric subduction is a new geodynamic model for this area, and it explains the tectonic, magmatic and metallogenic characteristics of the Lut Block.

Acknowledgements

The authors wish to thank Mrs. Sara Ribeiro (Laboratório de Geologia Isotópica da Universidade de Aveiro) for the TIMS analysis and for the guidance and assistance during sample preparation in the clean room. This research was financially supported by the Geobiotec Research Unit,

University of Aveiro, Portugal. Ministry of Sciences, Research and Technology of Iran is thanked for financial support for sabbatical research of Reza Arjmandzadeh in Portugal. National Iranian Copper Industries Company is thanked for the assistance with drill hole studies.

References

- [1] Stocklin J. *Structural history and tectonics of Iran: A review*-Amer. Ass. Petrol. Geol. Bull. 52 (1968) 1229-1258.
- [2] Berberian M., King G.C. "Towards a paleogeography and tectonic evolution of Iran". Canadian Journal of Earth Sciences. 18 (1981) 210-265.
- [3] Nabavi M.H. *An introduction to the geology of Iran*. Geological Survey of Iran. (1976) 109 p.
- [4] Tarkian M., Lotfi M., Baumann A. *Tectonic, magmatism and the formation of mineral deposits in the central Lut, east Iran*, Ministry of mines and metals, GSI, geodynamic project (geotraverse) in Iran. 51 (1983) 357-383.
- [5] Jung D., Keller J., Khorasani R., Marcks Chr., Baumann A., Horn P. *Petrology of the Tertiary magmatic activity the northern Lut area, East of Iran*, Ministry of mines and metals, GSI, geodynamic project (geotraverse) in Iran. 51 (1983) 285-336.
- [6] Samani B., Ashtari Sh. *Geological evolution of Sistan and Baluchestan area*, Journal of earth sciences. Geological Survey of Iran. No4 (1992).

Archive of SID

- [7] Saccani E., Delavari M., Beccaluva L., Amin S.A. *Petrological and geochemical constraints on the origin of the Nehbandan ophiolitic complex (eastern Iran): Implication for the evolution of the Sistan Ocean*. Lithos. (2010) Accepted Paper.
- [8] Eftekharnajad J. *Tectonic division of Iran with respect to sedimentary basins*. Journal of Iranian Petroleum Society. 82 (1981) 19–28 (in Persian).
- [9] Berberian M. *Continental deformation on the Iranian Plateau*, G.S.I. No. 52 (1983).
- [10] Tirrul R., Bell I.R., Griffis R.J., Camp V.E. *The Sistan suture zone of eastern Iran*. Geological Society of America Bulletin. 94 (1983) 134–150.
- [11] Doglioni C., Tonarini S., Innocenti F. *Mantle wedge asymmetries and geochemical signatures along W- and E-NE directed subduction zones*. Lithos. 113 (2009) 179–189.
- [12] Arjmandzadeh R., Karimpour M.H., Mazaheri S.A., Santos J.F., Medina J.M., Homam S.M. *Two sided asymmetric subduction: new hypothesis for the tectonomagmatic and metallogenic setting of the Lut Block, Eastern Iran*. 1st Symposium of Society of Economic Geology of Iran, Ferdowsi University of Mashhad. (2010).
- [13] Berberian M. *Active faulting and tectonics of Iran*, in Gupta, H. K., and Delany, F. M., editors, Zagros-Hindu Kush-Himalaya Geodynamic Evolution: American Geophysical Union Geodynamic Series. v. 3 (1981) 33–69.
- [14] Jackson J., McKenzie D. *Active tectonics of the Alpine-Himalayan Belt between western Turkey and Pakistan*: Geophysical Journal of the Royal Astronomical Society, v. 77 (1984) 185–264.
- [15] Lindenberg H.G., Großler K., Jacobshagen V., Ibbeken H. *Post-Paleozoic stratigraphy, structure and orogenetic evolution of the southern Sabzevar zone and the Taknar block*: Neues Jahrbuch für Geologie und Paläontologie, Abhandlungen, v. 168 (1984) 287–326.
- [16] Haghipour A., Aghanabati A. *Geological Map of Iran* (2nd edition): Tehran, Geological Survey of Iran, scale 1:2,500,000 (1989).
- [17] Alavi M. *Tectonic map of the Middle East*: Tehran, Geological Survey of Iran, scale 1:5,000,000 (1991).
- [18] Isacks B.L., Barazangi M. *Geometry of Benioff zones: lateral segmentation and downward bending of the subducted lithosphere*. In: Talwani, M., Pitman, W.M. III Eds., Island Arcs, Deep Sea Trenches and Back-arc Basins. AGU, Maurice Ewing Series 1 (1977) 99–114.
- [19] Doglioni C., Carminati E., Cuffaro M., Scrocca, D. *Subduction kinematics and dynamic constraints*. Earth Science Reviews. 83 (1977) 125–175. doi:10.1016/j.earscirev.2007.04.001.
- [20] Conder, J.A., Wiens D.A. *Seismic structure beneath the Tonga arc and Lau backarc basin determined from joint Vp, Vp/Vs tomography*. Geochemistry, Geophysics, and Geosystems 7, Q03018. (2006).
- [21] Heit B. *Teleseismic tomographic images of the Central Andes at 21°S and 25.5°S*. FU Berlin, Digitale Dissertation, <http://www.diss.fu-berlin.de/2005/319/>. (2005).
- [22] Tatsumi Y., Eggins S. *Subduction Zone Magmatism*. Blackwell Science, Cambridge, U.K. (1995) 211 pp.
- [23] Peccerillo A., Taylor S.R. *Geochemistry of Eocene calcalkaline volcanic rocks from the Kastamonu area, northern Turkey*. Contributions to Mineralogy and Petrology. 58 (1976) 63–81.
- [24] Pearce J.A., Harris, N.B.W., Tindle A.G. *Trace element discrimination diagrams for the tectonic interpretation of granitic rocks*. Journal of Petrology. 25 (1984) 956–983.
- [25] McDonough W.F., Frey F.A. *Rare earth elements in upper mantle rocks*. Mineralogical Society of America. Reviews in Mineralogy. (1989).
- [26] Wilson M. *Igneous Petrogenesis: A Global Tectonic Approach*. Harper Collins Academic. (1989) 466 p.
- [27] Martin H. *The adakitic magmas: modern analogues of Archaean granitoids*. Lithos. 46 (1999) 411–429.
- [28] Walker J.A., Patino, L.C., Carr, M.J., Feigenson, M.D. *Slab control over HFSE depletions in central Nicaragua*. Earth Planet. Sci. Lett. 192 (2001) 533–543.
- [29] Lan C.Y., Jahn B.M., Mertzman S.A., Wu, T.W. *Subduction-related granitic rocks of Taiwan*. Journal Southeast Asian Earth Science. 14 (1996) 11–28.
- [30] Kepezhinskas P.K., McDermott F., Defant M.J., Hochstaedter F.G., Drummond M.S., Hawkesworth C.J., Koloskov A., Maury R.C., Bellon H. *Trace element and Sr–Nd–Pb isotopic constraints on a three-component model of Kamchatka arc*

Archive of SID

petrogenesis. *Geochim. Cosmochim. Acta*. 61 (1997) 577–600.

[31] Castillo P.R., Janney P.E., Solidum R.U. *Petrology and geochemistry of Camiguin Island, southern Philippines: Insights to the source of adakites and other lavas in a complex arc setting: Contributions to Mineralogy and Petrology*. v. 134 (1999) 33–51.

[32] Defant M.J., Drummond M.S., *Derivation of some modern arc magmas by melting of young subducted lithosphere*. *Nature*. 347 (1990) 662–665.

[33] Grove T.L., Chatterjee N., Parman S.W., Médard E. *The influence of H₂O on mantle wedge melting: Earth and Planetary Science Letters*. 249 (2006) p. 74–89.

[34] Macpherson C.G., Dreher S.T., Thirlwall M.F. *Adakites without slab melting: High pressure differentiation of island arc magma, Mindanao, the Philippines: Earth and Planetary Science Letters*. 243 (2006) p. 581–593.

[35] Stern R.A., Hanson G.N. *Archean high-Mg granodiorite: a derivative of light rare earth element enriched monzodiorite of mantle origin*. *Journal of Petrology*. 32 (1991) 201–238.

[36] Wang Q., McDermott F., Xu J.F., Bellon H., Zhu Y.T. *Cenozoic K-rich adakitic volcanic rocks in the Hohxil area, northern Tibet: lower-crustal melting in an intracontinental setting*. *Geology*. 33 (2005) 465–468.

[37] Guo F., Fan W.M., Li C.W. *Geochemistry of late Mesozoic adakites from the Sulu belt, eastern China: magma genesis and implications for crustal recycling beneath continental collisional orogens*. *Geological Magazine*. 143 (2006) 1–13.

[38] Lai S.C., Qin J.F., Li Y.F. *Partial melting of thickened Tibetan Crust: geochemical evidence from Cenozoic adakitic volcanic rocks*. *International Geological Review*. 49 (2007) 357–373.

[39] Liu S., Hu R.Z., Feng C.X., Zou H.B., Li C., Chi X.G., Peng J.T., Zhong H., Qi L., Qi Y.Q., Wang T. *Cenozoic high Sr/Y volcanic rocks in the Qiangtang terrane, northern Tibet: geochemical and isotopic evidence for the origin of delaminated lower continental melts*. *Geological Magazine*. 145 (4) (2008a) 463–474.

[40] Liu S., Hu R.Z., Gao S., Feng C.X., Qi Y.Q., Wang T., Feng G.Y., Coulson, I.M. *U–Pb zircon age, geochemical and Sr–Nd–Pb–Hf isotopic constraints on age and origin of alkaline intrusions and*

associated mafic dikes from Sulu orogenic belt, Eastern China. *Lithos*. 106 (2008b) 365–379.

[41] Martin H. *Effect of steeper Archean geothermal gradient on geochemistry of subduction-zone magmas: Geology*. 14 (1986) p. 753–756.

[42] Rapp R.P., Watson E.B. *Dehydration melting of metabasalt at 8–32 kbar: Implications for continental growth and crust-mantle recycling: Journal of Petrology*. 36 (1995) 891–931.

[43] Liu Y.S., Gao S., Kelemen P.B., Xu W.L. *Recycled crust controls contrasting source compositions of Mesozoic and Cenozoic basalts in the North China Craton*. *Geochimica et Cosmochimica Acta*. 72 (2008) 2349–2376.

[44] Foley S., Tiepolo M., Vannucci R. *Growth of early continental crust controlled by melting of amphibolite in subduction zones: Nature*. 417 (2002) 837–840.

[45] Ionov D.A., Hofmann A.W. *Na–Ta-rich mantle amphiboles and micas: implications for subduction-related metasomatic trace element fractionations*. *Earth Planet. Sci. Lett.* 131 (1995) 341–356.

[46] Prouteau G., Scaillet B., Pichavant M., Maury R. *Evidence for mantle metasomatism by hydrous silicic melts derived from subducted oceanic crust*. *Nature*. 410 (2001) 197–200.

[47] Alirezai S., Arjmandzadeh R. *Mivehrood adakitic porphyry, NW Iran; Tectonic and economic implications*. Congress of GAC-MAC; Annual meeting, Montréal (2006).

[48] Aftabi A., Atapour H. *Regional aspects of shoshonitic volcanism in Iran*. *Episodes*. 23 (2000) no. 2, 119–125.

[49] Shahabpour J. *Economic Geology*. Publication of Shahid Bahonar University of Kerman. (2001) 543p.

[50] Karimpour M.H., Stern C.R. *Advanced spaceborne thermal emission and reflection radiometer mineral mapping to discriminate high sulfidation, reduced intrusion related and iron oxide gold deposits, Eastern Iran*. *Journal of Applied Sciences*. 9 (5) (2009) 815–828.

[51] Sengör A.M.C., Altner D., Cin, A., Ustaomer T., Hsu K.J. *Origin and assembly of the Tethyside orogenic collage at the expense of Gondwana Land*, In: Audley-Charles, M.G., Hallam, A.E. (Eds.), *Gondwana and Tethys*. Geological Society of London Special Publication, Blackwell, Oxford. (1988) 119–181.

Archive of SID

- [52] Camp V.E., Griffis R.J. *Character, genesis and tectonic setting of igneous rocks in the Sistan suture zone, eastern Iran*. Lithos. 15 (1982) 221-239.
- [53] Berberian M., Yeats R.S. *Patterns of historical earthquake rupture in the Iranian Plateau*. Bulletin of the society of America. 89 (1999) 120-139.
- [54] Malekzadeh A. *Geology, mineralization, alteration, geochemistry, microthermometry, isotope*

- studies and determining the mineralization source of Khoopic and Maherabad exploration areas*. Ph.D thesis. Ferdowsi University of Mashhad. (2009).
- [55] Sengör A.M.C., Natalin B.A. *Paleotectonics of Asia: fragment of a synthesis*. In: An Y, Harrison TM (eds) *The tectonic evolution of Asia*. Cambridge Univ. Press, Cambridge. (1996) pp 486–640.