

Linear Stability Analysis of Double Diffusive Convection in a Horizontal Sparsely Packed Rotating Anisotropic Porous Layer in Presence of Soret Effect

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ABSTRACT

In this paper, we have investigated theoretically the effect of Soret parameter on the onset of double diffusive rotating anisotropic convection in a horizontal sparsely packed porous layer using linear stability theory which is based on the usual normal mode technique. The Brinkman model that includes the Coriolis term is employed for the momentum equation. The effect of anisotropy parameters, Soret parameter, solute Rayleigh number, Taylor number, Lewis number, Darcy and Darcy Prandtl number on stationary and oscillatory convection is shown graphically.

Keywords: Soret parameter, Rotation, Anisotropy, Double diffusive convection.

NOMENCLATURE

a	wave number	ε	porosity
c	specific heat of solid	ρ_0	reference density
c_p	specific heat of fluid at constant pressure	η	thermal anisotropy parameter, D_{1x}/D_{1z}
d	height of the porous layer	μ_f	fluid viscosity
D_1	anisotropic thermal diffusion tensor, $D_{1x}\mathbf{ii} + D_{1y}\mathbf{jj} + D_{1z}\mathbf{kk}$	μ_e	effective viscosity
D_{11}	soret coefficient	ν	kinematic viscosity
D_2	solute Diffusivity	Θ	dimensionless amplitude of temperature perturbation
Da	Modified Darcy number, $\mu_e K_z / \mu_f d^2$	σ	growth rate
\mathbf{g}	gravitational acceleration, $(0, 0, -g)$	Ω	angular velocity of rotation, $(0, 0, \Omega)$
K	inverse anisotropic permeability tensor, $K_x^{-1}\mathbf{ii} + K_y^{-1}\mathbf{jj} + K_z^{-1}\mathbf{kk}$	ξ	mechanical anisotropy parameter, K_x/K_z
l, m	horizontal wavenumbers	λ	normalized porosity, ε/γ
Le	Lewis number, D_{1z}/D_2	Φ	dimensionless amplitude of concentration perturbation
p	pressure	ψ	stream function
Pr_D	Darcy-Prandtl number, $\gamma \varepsilon \nu d^2 / K_z D_{1z}$	Other symbols	
\mathbf{q}	velocity vector, (u, v, w)	∇_h^2	$\frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial y^2}$
R_T	thermal Rayleigh number, $\beta_T g \Delta T d K_z / \nu D_{1z}$	∇^2	$\nabla_h^2 + \frac{\partial^2}{\partial z^2}$
R_S	solute Rayleigh number, $\beta_S g \Delta S d K_z / \nu D_{1z}$	Subscripts	
		b	basic state

S	solute concentration
Sr	soret parameter, $D_{11}\beta_S / D_{1z} \beta_T$
t	time
T	temperature
Ta	Taylor number, $(2\Omega K_Z / \varepsilon\nu)^2$
ΔS	salinity difference between the walls
ΔT	temperature difference between the walls
x, y, z	space coordinates

Greek symbols

β_T	thermal expansion coefficient
β_S	solute expansion coefficient
γ	ratio of specific heat, $\frac{(\rho c)_m}{(\rho c)_f}$

c	critical
f	fluid
h	horizontal
s	solid
m	porous medium
0	reference value

Superscripts

*	dimensionless quantity
'	perturbed quantity
<i>osc</i>	oscillatory state
<i>st</i>	stationary state

1. INTRODUCTION

A double diffusive process is one that is driven by the presence of two different substances that diffuse at different rates. Double diffusion was first explored because of its application to oceanic phenomena. In the ocean, the density of water is governed primarily by temperature and salinity. The applications of the present study has recently increased in view of its potential occurrence in nature and wide range of applications such as high-quality crystal production, liquid gas storage, migration of moisture in fibrous insulation, transport of contaminants in saturated soil, solidification of molten alloys, and geothermally heated lakes and magmas, underground disposal of nuclear wastes, liquid re-injection, electro-chemical and drying processes.

A comprehensive review of the literature concerning double diffusive natural convection in a fluid-saturated porous medium may be found in the book by [Nield and Bejan \(2006\)](#). Useful review articles on double diffusive convection in porous media include those by [Trevisan and Bejan \(1999\)](#), [Mojtabi and Charrier-Mojtabi \(2005\)](#) and [Mamou \(2002\)](#), [Vafai \(2005\)](#) and [Vadasz \(2008\)](#), [Ingham and Pop \(2005\)](#).

The onset of double diffusive convection in a horizontal porous layer has been investigated by [Rudraiah et al. \(1982\)](#) using a weak non-linear theory. Finite amplitude double diffusive convection near the threshold of both stationary and oscillatory instabilities in a binary mixture was investigated by [Brand and Steinberg \(1983\)](#). The linear stability analysis of the thermosolutal convection in a sparsely packed porous layer was made by [Poulikakos \(1986\)](#) using the Darcy-Brinkman model. Small amplitude nonlinear solutions in the form of standing and traveling waves and the transition to finite amplitude convection, as predicted by bifurcation theory, were studied by [Knobloch \(1986\)](#). The double diffusive convection in porous media in the presence of Soret

and Dufour coefficients has been analyzed by [Rudraiah and Malashetty \(1986\)](#).

[Murray and Chen \(1989\)](#) have extended the linear stability theory, taking into account effects of temperature-dependent viscosity and volumetric expansion coefficients and nonlinear basic salinity profile. Double diffusive fingering convection in a porous medium with horizontally periodic boundary conditions was studied by [Chen and Chen \(1993\)](#). [Malashetty \(1993\)](#) made a linear stability analysis to determine the effects of anisotropic thermo convective currents on the double diffusive convection in a sparsely packed porous medium.

[Straughan and Hutter \(1999\)](#) have investigated the double diffusive convection with Soret effect in a porous layer using Darcy-Brinkman model and derived a priori bounds. An analytical and numerical study of double diffusive convection with parallel flow in a horizontal sparsely packed porous layer under the influence of constant heat and mass flux was performed using a Brinkman model by [Amahmid et al. \(1999\)](#).

[Bahloul et al. \(2003\)](#) have carried out an analytical and numerical study of the double diffusive convection in a shallow horizontal porous layer under the influence of Soret effect. [Hill \(2005\)](#) performed linear and nonlinear stability analyses of double diffusive convection in a fluid saturated porous layer with a concentration based internal heat source using Darcy's law. Double diffusive natural convection within a multilayer anisotropic porous medium is studied numerically and analytically by [Bennacer et al. \(2005\)](#). [Mansour et al. \(2006\)](#) have investigated the multiplicity of solutions induced by thermosolutal convection in a square porous cavity heated from below and subjected to horizontal concentration gradient in the presence of Soret effect.

There are only few studies available on double diffusive convection in the presence of rotation. [Chakrabarti and Gupta \(1981\)](#) have analyzed the nonlinear thermohaline convection in a rotating porous medium. The effect of rotation on linear and

non-linear double diffusive convection in a sparsely packed porous medium was studied by Rudraiah *et al.* (1986).

Anisotropy is generally a consequence of preferential orientation or asymmetric geometry of porous matrix or fibers and is in fact encountered in numerous systems in industry and nature. Anisotropy is particularly important in a geological context, since sedimentary rocks generally have a layered structure; the permeability in the vertical direction is often much less than in the horizontal direction. The review of research on convective flow through anisotropic porous media has been well documented by McKibbin (1992) and Storesletten (2004). Castinel and Combarnous (1974) have conducted an experimental and theoretical investigation on the Rayleigh-Benard convection in an anisotropic porous medium. Epherre (1977) extended the stability analysis to a porous medium with anisotropy in thermal diffusivity also.

Recently many authors have studied the effect of anisotropy and/or rotation on the onset of convection in a horizontal porous layer (see e.g. Govender; 2007, Malashetty and Swamy; 2007). There appears to be only couple of studies available on the double diffusive convection in an anisotropic porous medium with rotation (Patil *et al.*, 1989) and without rotation (Tyvand; 1980).

Recently Gaikwad and Kamble (2012) have investigated the Soret effect on double diffusive convection in a horizontal sparsely packed porous layer. Rani and Reddy (2013) have studied the Soret and Dufour Effects on Transient Double Diffusive Free Convection of Couple-stress Fluid past a Vertical Cylinder. The effect of rotation on the onset of double diffusive convection in a sparsely packed anisotropic porous layer is studied by Malashetty and Begum (2011). A very little attention has been devoted to the study of double diffusive convection in rotating porous medium including the mechanical and thermal anisotropic effects in presence of Soret effect. Therefore, the objective of the present study is to investigate the combined effect of rotation and anisotropy in presence of Soret effect on the double diffusive convection in a horizontal sparsely packed porous layer using linear stability analyses.

2. GOVERNING EQUATIONS

Consider a horizontal sparsely packed anisotropic porous layer, saturated with Boussinesq fluid of infinite horizontal extent confined between parallel, stress-free planes $z = 0$ and $z = d$ subject to rotation and maintained at constant temperatures $T_0 + \Delta T$ and T_0 with solute concentration $S_0 + \Delta S$ and S_0 respectively. A Cartesian frame of reference is chosen with x -and y -axes at the lower boundary plane and z -axis directed vertically upwards in the gravity field. The axis of rotation is assumed to coincide with the z -axis with a constant angular velocity $\Omega = (0, 0, \Omega)$. The

Darcy-Brinkman model that includes the Coriolis term is used for the momentum equation. The velocities are assumed to be small so that the advective and Forchheimer inertia effects are ignored. The Boussinesq approximation, which states that the variation in density is negligible every where in the conservations except in the buoyancy term, is assumed to hold. With these assumptions the basic governing equations are

$$\nabla \cdot \mathbf{q} = 0 \tag{1}$$

$$\rho_0 \left(\frac{1}{\varepsilon} \frac{\partial \mathbf{q}}{\partial t} + \frac{2}{\varepsilon} \boldsymbol{\Omega} \times \mathbf{q} \right) + \mu_f \mathbf{K} \cdot \mathbf{q} = -\nabla p + \rho \mathbf{g} + \mu_e \nabla^2 \mathbf{q} \tag{2}$$

$$\gamma \frac{\partial T}{\partial t} + (\mathbf{q} \cdot \nabla) T = \nabla \cdot (D_1 \cdot \nabla T) \tag{3}$$

$$\varepsilon \frac{\partial S}{\partial t} + (\mathbf{q} \cdot \nabla) S = D_2 \nabla^2 S + D_{11} \nabla^2 T \tag{4}$$

$$\rho = \rho_0 (1 - \beta_T (T - T_0) + \beta_S (S - S_0)) \tag{5}$$

The basic state of the fluid is assumed to be quiescent. The quantities of the basic state are given by

$$\mathbf{q}_b = (0, 0, 0), p = p_b(z), T = T_b(z), S = S_b(z), \rho = \rho_b(z) \tag{6}$$

which satisfy the equations

$$\frac{dp_b}{dz} = -\rho_b g, \frac{d^2 T_b}{dz^2} = 0, \frac{d^2 S_b}{dz^2} = 0, \rho_b = \rho_0 (1 - \beta_T (T_b - T_0) + \beta_S (S_b - S_0)) \tag{7}$$

On the basic state we superpose small perturbations around the basic solutions in the form

$$\mathbf{q} = \mathbf{q}_b + \mathbf{q}', T = T_b + T', S = S_b + S', \rho = \rho_b + \rho' \tag{8}$$

where the primes indicate perturbations.

Substituting Eq. (8) into Eqs. (1) to (5) and using the basic state Eqs. (7) and transformations

$$(x', y', z') = (x^*, y^*, z^*)d, t' = t^* (\gamma d^2 / D_{1z}),$$

$$(u', v', w') = (D_{1z} / d) (u^*, v^*, w^*),$$

$$T' = (\Delta T) T^*, S' = (\Delta S) S^* \tag{9}$$

to render the equations dimensionless we obtain (after dropping the asterisks for simplicity

$$w \left[\frac{1}{Pr_D} \frac{\partial}{\partial t} \nabla^2 + \frac{1}{\xi} \frac{\partial^2}{\partial z^2} - Da \nabla^4 + \nabla_h^2 \right] \times \left[Ta \frac{\partial^2}{\partial z^2} + \left(\frac{1}{Pr_D} \frac{\partial}{\partial t} + \frac{1}{\xi} - Da \nabla^2 \right) \right] = \left(\frac{1}{Pr_D} \frac{\partial}{\partial t} + \frac{1}{\xi} - Da \nabla^2 \right) (R_T \nabla_h^2 T - R_S \nabla_h^2 S) \tag{10}$$

$$\left[\frac{\partial}{\partial t} - \left(\eta \nabla_h^2 + \frac{\partial^2}{\partial z^2} \right) + \mathbf{q} \cdot \nabla \right] T - w = 0 \quad (11)$$

$$\left[\lambda \frac{\partial}{\partial t} - \frac{1}{Le} \nabla^2 + \mathbf{q} \cdot \nabla \right] S - w = Sr \frac{R_T}{R_S} \nabla^2 T \quad (12)$$

Where $Pr_D = \gamma \epsilon \nu d^2 / K_z D_{1z}$, the Darcy-Prandtl number, $Da = \mu_e K_Z / \mu_f d^2$, the modified Darcy number, $Ta = (2\Omega K_Z / \epsilon \nu)^2$, the Taylor number, $Sr = D_{11} \beta_S / D_{1z} \beta_T$, the Soret parameter $R_T = \beta_T g \Delta T d K_z / \nu D_{1z}$, the thermal Rayleigh number, $R_S = \beta_S g \Delta S d K_z / \nu D_{1z}$, the solute Rayleigh number, $Le = D_{1z} / D_2$, the Lewis number, $\xi = K_x / K_z$, the mechanical anisotropy parameter, $\eta = D_{1x} / D_{1z}$, the thermal anisotropy parameter, $\lambda = \epsilon / \gamma$, normalized porosity. The boundary conditions in only the z -direction are required to solving Eqs. (10) to (12) and are given by

$$w = \frac{\partial^2 w}{\partial z^2} = T = S = 0 \quad \text{at } z = 0 \quad (13)$$

3. LINEAR STABILITY ANALYSIS

In this section we predict the thresholds of both , stationary and oscillatory convections using linear theory. The Eigen value problem defined by Eqs. (10) to (12) subject to the boundary conditions (13) is solved using the time-dependent periodic disturbances in a horizontal plane. Assuming that the amplitudes of the perturbations are very small, we write

$$\begin{pmatrix} w \\ T \\ S \end{pmatrix} = \begin{pmatrix} W(z) \\ \Theta(z) \\ \Phi(z) \end{pmatrix} \exp [i (lx + my) + \sigma t] \quad (14)$$

where l, m are horizontal wavenumbers and σ is the growth rate. Infinitesimal perturbations of the rest state may either damp or grow depending on the value of the parameter σ . Substituting Eq. (14) into the linearized version of Eqs. (10) to (12) we obtain

$$w \begin{pmatrix} Ta D^2 + \left(\frac{\sigma}{Pr_D} (D^2 - a^2) - \left((D^2 - a^2)^2 Da + \frac{1}{\xi} D^2 - a^2 \right) \right) \\ \left(\frac{\sigma}{Pr_D} + \frac{1}{\xi} - (D^2 - a^2) Da \right) \\ \left(\frac{\sigma}{Pr_D} + \frac{1}{\xi} - (D^2 - a^2) Da \right) \\ \left(R_S a^2 \Phi - R_T a^2 \Theta \right) \end{pmatrix} = \begin{pmatrix} \left(\frac{\sigma}{Pr_D} + \frac{1}{\xi} - (D^2 - a^2) Da \right) \\ \left(R_S a^2 \Phi - R_T a^2 \Theta \right) \end{pmatrix} \quad (15)$$

$$\left[\sigma - (D^2 - a^2 \eta) \right] \Theta - W = 0 \quad (16)$$

$$\left[\sigma \lambda - \frac{1}{Le} (D^2 - a^2) \right] \Phi - W = B (D^2 - a^2) \Theta \quad (17)$$

Where, $B = Sr \frac{R_T}{R_S}$, $D = d/dz$ and $a^2 = l^2 + m^2$.

We assume the solutions of Eqs. (15) to (17) satisfying the boundary conditions (13) in the form

$$\begin{pmatrix} W(z) \\ \Theta(z) \\ \Phi(z) \end{pmatrix} = \begin{pmatrix} W_0 \\ \Theta_0 \\ \Phi_0 \end{pmatrix} \sin n \pi z, (n = 1, 2, 3, \dots) \quad (18)$$

The most unstable mode corresponds to $n = 1$ (fundamental mode). Therefore substituting Eq. (18) with $n = 1$ into Eqs. (15) to (17), we obtain a matrix equation

$$\begin{pmatrix} A_{11} & -a^2 R_T & a^2 R_S \\ -1 & \sigma + \delta_2^2 & 0 \\ -1 & B \delta^2 & \sigma \lambda + \delta^2 Le^{-1} \end{pmatrix} \begin{pmatrix} W_0 \\ \Theta_0 \\ \Phi_0 \end{pmatrix} = \begin{pmatrix} 0 \\ 0 \\ 0 \end{pmatrix} \quad (19)$$

For brevity $A_{11}, \delta^2, \delta_1^2$ and δ_2^2 are not mentioned here. The condition of nontrivial solution of above system of homogeneous linear Eqs (19) yields the expression for thermal Rayleigh number in the form

$$R_T = \left\{ \frac{\left[\left(\frac{\sigma}{Pr_D} + \delta^2 Da \right) \delta^2 + \frac{Ta \pi^2}{\left(\sigma Pr_D^{-1} + \frac{1}{\xi} + \delta^2 Da \right)} + \delta_1^2 \right]}{\left(\lambda \sigma + \delta^2 Le^{-1} \right) + \left(\sigma \lambda + (Le^{-1} + Sr) \delta^2 \right)} \right\} \left\{ \frac{(\sigma + \delta_2^2)}{a^2} \right\} \quad (20)$$

3.1 Stationary State

If σ is real, then marginal stability occurs when $\sigma = 0$ (i.e. $\omega_r = \omega_i = 0$). Then the Rayleigh number R_T^{St} for the onset of stationary convection, in the form

$$R_T^{St} = \frac{(a^2\eta + \pi^2)}{a^2(1+SrLe)} \left[\begin{aligned} & \left(\frac{a^2 + \pi^2 \xi^{-1}}{a^2 + \pi^2} \right) + \\ & \left(\frac{a^2 + \pi^2}{a^2 + \pi^2} \right)^2 Da \\ & + \frac{Ta\pi^2}{\left(\xi^{-1} + (a^2 + \pi^2)Da \right)} \end{aligned} \right] \\ \times \frac{(a^2\eta + \pi^2)}{(1+SrLe)(a^2 + \pi^2)} R_S Le. \tag{21}$$

In the absence of Soret effect i.e., $Sr = 0$ Eq. (21) gives

$$R_T^{St} = \frac{(a^2\eta + \pi^2)}{a^2} \left[\begin{aligned} & \left(a^2 + \frac{\pi^2}{\xi} \right) + (a^2 + \pi^2)^2 Da \\ & + \frac{Ta\pi^2}{\frac{1}{\xi} + (a^2 + \pi^2)Da} \end{aligned} \right] \\ \frac{(a^2\eta + \pi^2)}{(a^2 + \pi^2)} R_S Le. \tag{22}$$

The minimum value of the Rayleigh number R_T^{St}

occurs at the critical wavenumber $a = a_c^{St}$. In the limit as $Da \rightarrow 0$ i.e., for a densely packed porous medium Eq. (22) reduces to

$$R_T^{St} = \frac{(a^2\eta + \pi^2) \left[\left(a^2 + \frac{\pi^2}{\xi} \right) + \pi^2 Ta\xi \right]}{a^2} \\ + \left(\frac{a^2\eta + \pi^2}{a^2 + \pi^2} \right) R_S Le. \tag{23}$$

This is exactly the one given by Malashatty and Heera (2008). When $Da \rightarrow 0$ and $Ta=0$, i.e. for a densely packed porous medium in the absence of rotation, Eq. (23) reduces to

$$R_T^{St} = \frac{1}{a^2} \left(a^2 + \frac{\pi^2}{\xi} \right) (a^2\eta + \pi^2) \\ + \left(\frac{a^2\eta + \pi^2}{a^2 + \pi^2} \right) R_S Le \tag{24}$$

given by Malashatty and Swamy (2009). Further, for an isotropic porous media, that is when $\xi = \eta = 1$, Eq. (22) gives

$$R_T^{St} = \frac{\pi^4 (1 + \alpha^2)^3 M}{\alpha^2} + \frac{Ta}{M\alpha^2} + \frac{R_S}{\tau}, \tag{25}$$

Where

$$M = \frac{1}{\pi^2} \left(Da\pi^2 + \left(1 + \frac{a^2}{\pi^2} \right)^{-1} \right), \quad \alpha^2 = \frac{a^2}{\pi^2}, \quad \tau = \frac{1}{Le},$$

which is the one obtained by Rudraiah et al. (1986).

3.2 Oscillatory State

We now set $\sigma = i\omega_i$ in Eq. (20) and clear the complex quantities from the denominator, to obtain

$$R_T = \Delta_1 + i\omega_i \Delta_2 \tag{26}$$

Since R_T is a physical quantity, it must be real. Hence, from Eq. (26) it follows that either $\omega_i = 0$ (steady onset) or $\Delta_2 = 0$ ($\omega_i \neq 0$, oscillatory onset).

For oscillatory onset $\Delta_2 = 0$ ($\omega_i \neq 0$) and this gives a dispersion relation of the form (on dropping the subscript i)

$$a_0 (\omega^2)^2 + a_1 (\omega^2) + a_2 = 0 \tag{27}$$

for simplicity the coefficients a_0, a_1 and a_2 are not mentioned here. Now Eq. (26) with $\Delta_2 = 0$, gives

$$R_T^{osc} = \frac{\left(\begin{aligned} & \left(\begin{aligned} & \omega^2 \left(\frac{Pr_D r_1 \delta^4 \xi + \lambda Le^2 \times}{\left(\omega^2 \xi \lambda - Pr_D^2 r_2 Sr \delta^2 \right)} \right) + \\ & \left(\begin{aligned} & \omega^2 \delta^2 Le Pr_D r_1 \lambda \xi \\ & - \omega^2 Le \delta^2 \lambda \xi + Pr_D^2 \times \\ & \left(\delta^4 r_1 + \omega^2 Le^2 \lambda^2 \right) r_2 \end{aligned} \right) \end{aligned} \right) \pi^2 \end{aligned} \right)}{\left(a^2 (\delta^4 r_1^2 + \omega^2 Le^2 \lambda^2) (Pr_D^2 r_2^2 + \omega^2 \xi^2) \right)} + \\ \frac{\left(\begin{aligned} & -\omega \delta^2 \left(\begin{aligned} & \omega^2 Le^2 \lambda^2 + \delta^4 (r_1 (1 + Da Le Pr_D \lambda)) \\ & + Le^2 Pr_D Sr \lambda \delta_1^2 \end{aligned} \right) + \\ & \left(\begin{aligned} & \delta^4 \left(\frac{-\omega^2 Le^2 Sr \lambda + Da Pr_D \times}{\left(\delta^4 r_1 + \omega^2 Le^2 \lambda^2 \right)} \right) + \delta_2^2 \\ & \left(Pr_D (\delta^4 r_1 + \omega^2 Le^2 \lambda^2) \delta_1^2 \right) \end{aligned} \right) \end{aligned} \right)}{\left(a^2 (\delta^4 r_1^2 + \omega^2 Le^2 \lambda^2) Pr_D \right)} \\ + \frac{Le R_S (\delta_2^2 \delta^2 r_1 + \omega^2 Le \lambda)}{\left(\delta^4 r_1^2 + \omega^2 Le^2 \lambda^2 \right)} \tag{28}$$

4. RESULT AND DISCUSSION

The Soret effect on the onset of double-diffusive rotating anisotropic convection in a horizontal sparsely packed porous layer is investigated theoretically using the linear stability theory. The variation of stationary and oscillatory curves in the $R_T - a$ plane for various parameter values are as shown in Figs.(1a to 4d). We fixed the values for the parameters except the varying parameter.

Figs (1a to 1d) and (3a to 3d) shows that the stationary and oscillatory curves respectively for different values of ξ, η, Ta and Da for fixed values of other parameters. We observed from these figures that the convection sets in as oscillatory mode prior to the stationary mode.

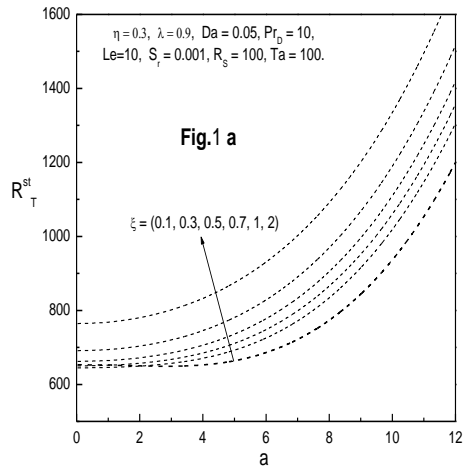


Fig. 1a. Variation of stationary curves for different values of ξ

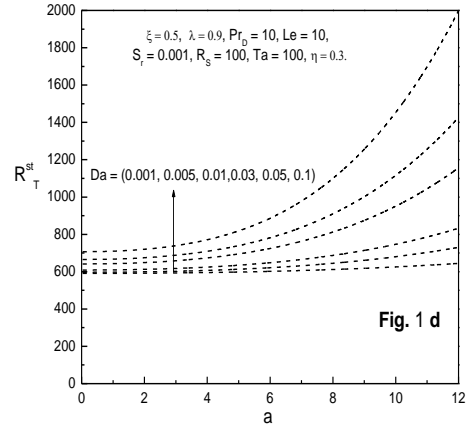


Fig. 1d. Variation of stationary curves for different values of Da

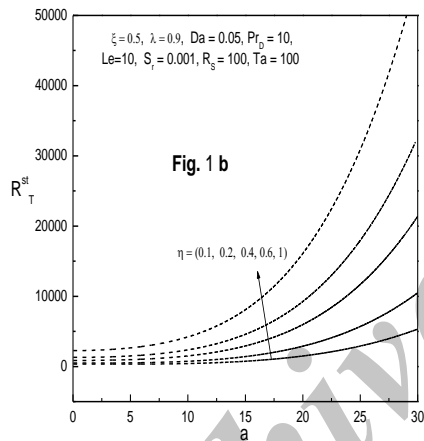


Fig. 1b. Variation of stationary curves for different values of η

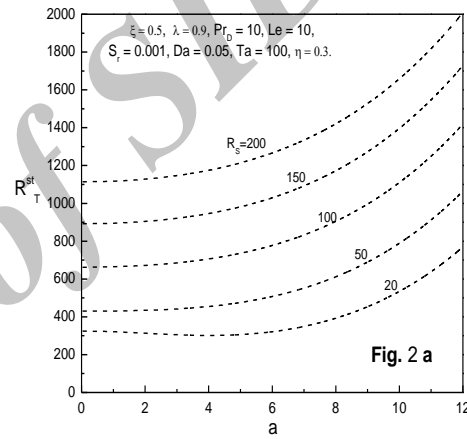


Fig. 2a. Variation of stationary curves for different values of R_s

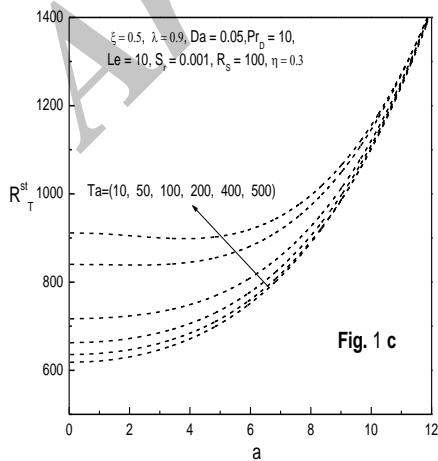


Fig. 1c. Variation of stationary curves for different values of Ta

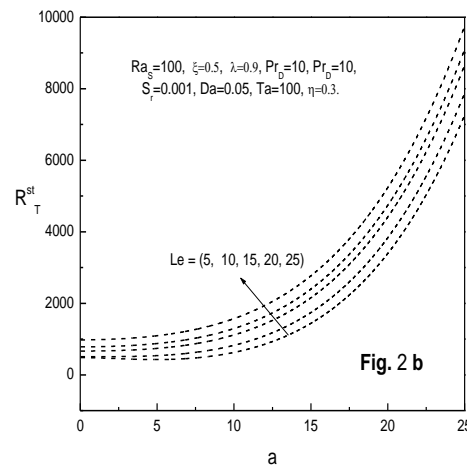


Fig. 2b. Variation of stationary curves for different values of Le

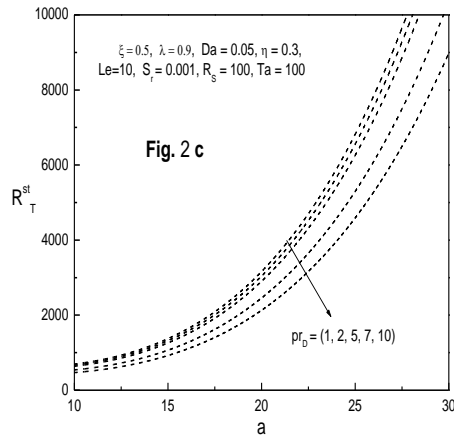


Fig. 2c. Variation of stationary curves for different values of Pr_D

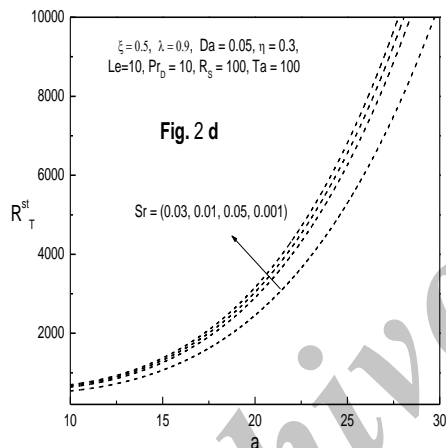


Fig. 2d. Variation of stationary curves for different values of Sr

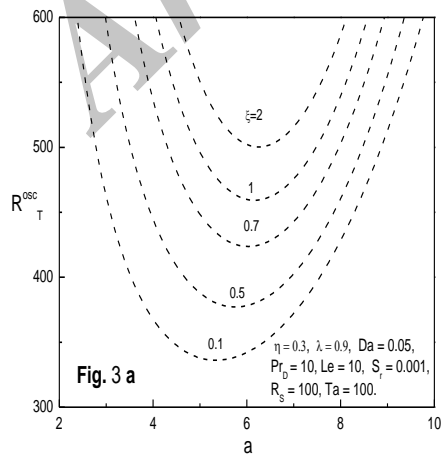


Fig. 3a. Variation of oscillatory curves for different values of ξ

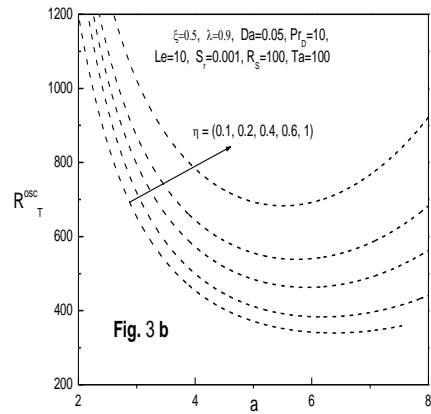


Fig. 3b. Variation of oscillatory curves for different values of η

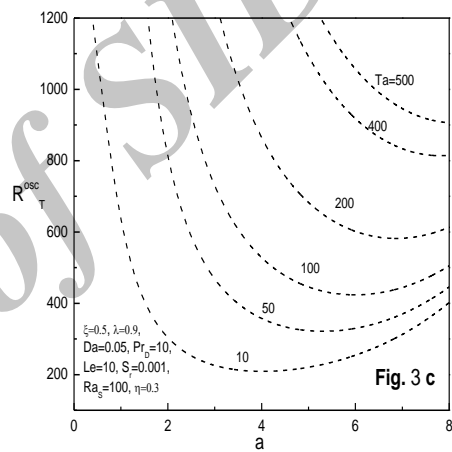


Fig. 3c. Variation of oscillatory curves for different values of Ta

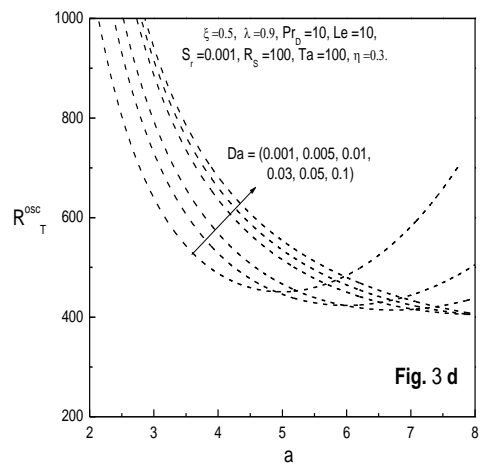


Fig. 3d. Variation of oscillatory curves for different values of Da

F

From Figs. 1a and 3a, it can be observed that the critical value of the Rayleigh number increases with

the increasing ξ for both oscillatory and stationary convection. Thus ξ has stabilizing effect on stationary and oscillatory convection. Figs. 1b and 3b indicates the effect of thermal anisotropy parameter η on both the stationary and oscillatory convection for the fixed values of other parameters. It is observed that critical value of the Rayleigh number for stationary and oscillatory mode increases with increasing η , indicating that the effect of η is to inhibit the onset of stationary and oscillatory convection. Figs. 1c and 3c depicts the effect of Taylor number Ta on the both the stationary and oscillatory curves. We find that the effect of increasing Ta is to increase the critical value of the Rayleigh number for stationary and oscillatory modes and the corresponding wavenumber. Thus the Taylor number Ta has a stabilizing effect on the double diffusive convection in a horizontal sparsely packed anisotropic porous medium. Figs. 1d and 3d present the effect of Darcy number Da on both the curves. We find that critical value of the stationary Rayleigh number increases with Da , indicating that the effect of Darcy number Da is to inhibit the onset of stationary convection. Whereas for the oscillatory convection, the critical value of oscillatory Rayleigh number decreases with increasing Da , and with further increase in the value of Da , the critical Rayleigh number increases. Thus Da has dual effect on oscillatory convection.

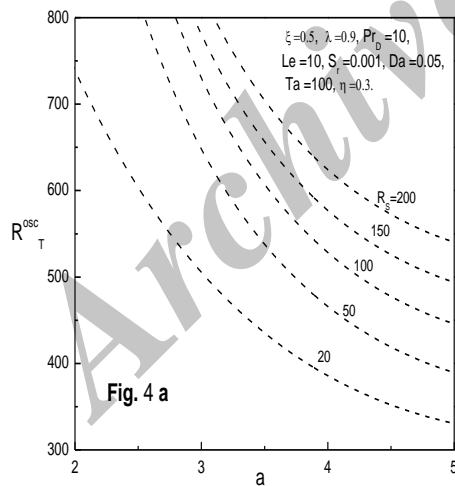


Fig. 4a. Variation of oscillatory curves for different values of R_S

Figs. (2a to 2d) and (4a to 4d) depict the effect of R_S , Le , Pr_D and Sr on stationary and oscillatory modes respectively. We find from Figs. 2a and 4a that the effect of increasing R_S is to increase the critical value of the Rayleigh number for stationary and oscillatory modes and the corresponding wavenumber. Thus the solute Rayleigh number R_S has a stabilizing effect on the

double diffusive convection in a horizontal sparsely packed anisotropic porous medium.

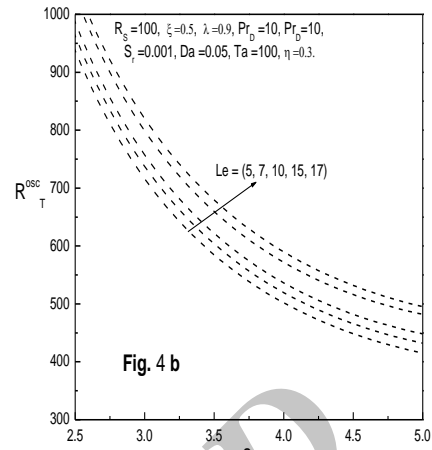


Fig. 4b. Variation of oscillatory curves for different values of Le

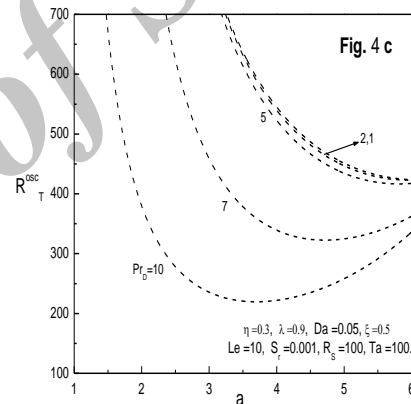


Fig.4 c. Variation of oscillatory curves for different values of Pr_D

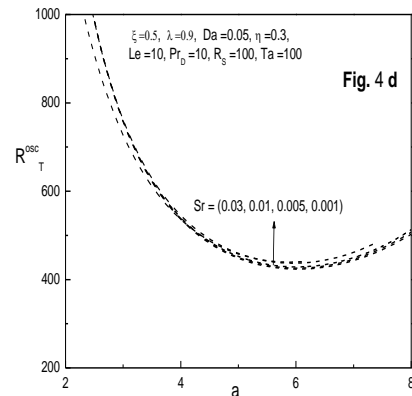


Fig. 4d. Variation of oscillatory curves for different values of Sr

In Figs. 2b and 4b the marginal and oscillatory curves for different values of Lewis number Le are drawn. It is observed that with the increase of Le the critical values of Rayleigh number and the corresponding wavenumber for oscillatory mode decreases while those for stationary mode increases. Therefore, the effect of Le is to advance the onset of oscillatory convection while its effect is to inhibit the stationary convection. The marginal and oscillatory curves for different values of Darcy-Prandtl number Pr_D are presented in Figs. 2c and 4c, from these figures it is evident that for small and moderate values of Pr_D the critical value of oscillatory Rayleigh number decreases with the increase of Pr_D , however this trend is reversed for large values of Pr_D . Figures. 2d and 4d indicates the effect of Soret parameter Sr on the both the stationary and oscillatory modes. It is observed that as Sr increases positively, the critical values of Rayleigh number and the corresponding wavenumber for oscillatory mode increase while as Sr increases negatively, those decrease. Whereas the effect is reversed for the stationary mode. Thus Soret parameter has stabilizing effect on oscillatory convection and destabilizing effect on stationary convection.

The detailed behavior of stationary and oscillatory critical Rayleigh number with respect to the Taylor number is analyzed in the $R_{Tc} - Ta$ plane through Figs. (5a to 5f) and (6a to 6f) respectively. We observe from these figures that the critical Rayleigh number increases with the increase of Ta , indicating that the effect of rotation is to inhibit the onset of thermal convection and it is in agreement with the corresponding problem of isotropic and pure fluid layer.

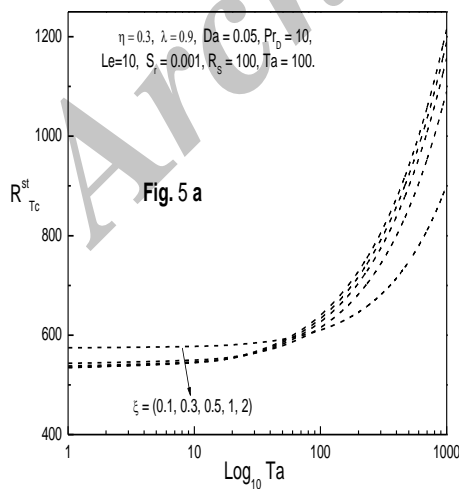


Fig. 5a. Variation of stationary critical curves for different values of ξ

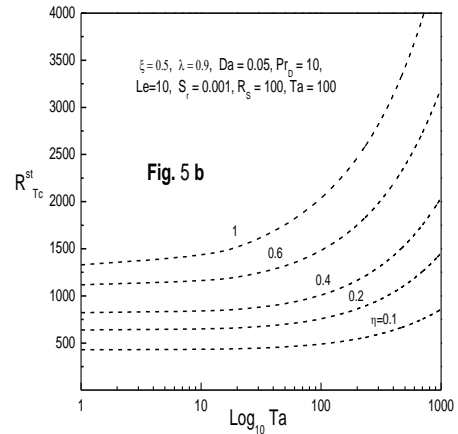


Fig. 5b. Variation of stationary critical curves for different values of η

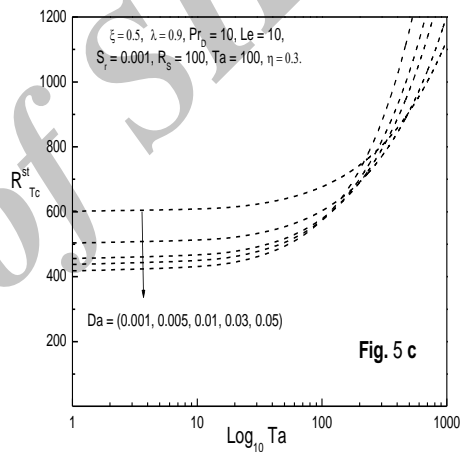


Fig. 5c. Variation of stationary critical curves for different values of Da

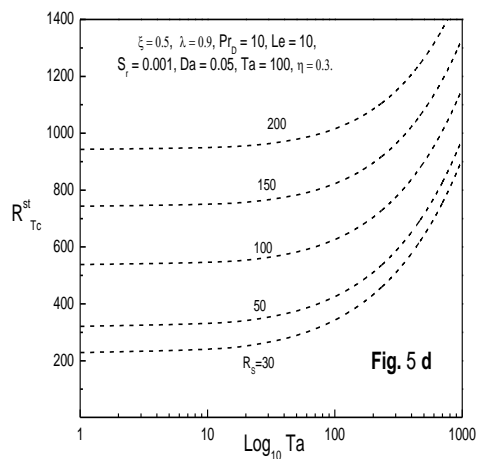


Fig. 5d. Variation of stationary critical curves for different values of R_s

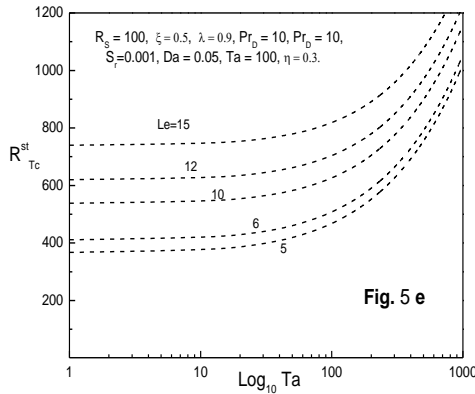


Fig. 5e. Variation of stationary critical curves for different values of Le

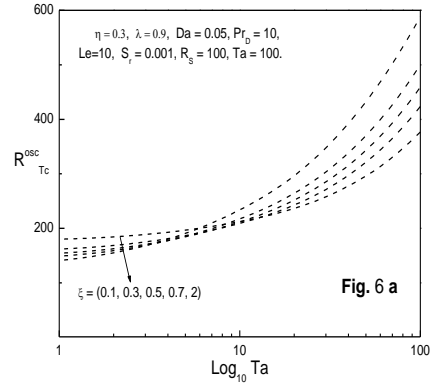


Fig. 6a. Variation of oscillatory critical curves for different values of ξ

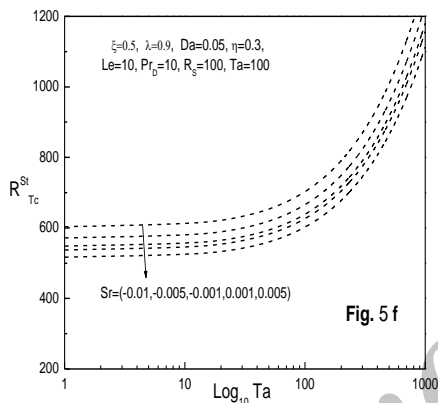


Fig. 5f. Variation of stationary critical curves for different values of Sr

In Figs. 5a and 6a, we display the variation of stationary and oscillatory critical Rayleigh number R_{Tc} with Taylor number Ta for different values of mechanical anisotropy parameter ξ for the fixed values of other parameters. It is important to note that R_{Tc} decreases with the increase of ξ for small values of Ta , and for large value of Ta , the critical Rayleigh number increases. Thus the mechanical anisotropy parameter ξ has dual effect on oscillatory and stationary convection. Figures. 5b and 6b indicates the variation of R_{Tc} with Ta for different values of thermal anisotropy parameter η . It is observed that the critical Rayleigh number R_{Tc} increases with the increase of η indicating that the effect of thermal anisotropy parameter is to inhibit the onset of stationary and oscillatory convection.

Figures. 5c and 6c presents the variation of R_{Tc} with Ta for different values of Darcy number Da . We find that the critical Rayleigh number R_{Tc} increases with the increase of Da for small values of Ta , and for large value of Ta , the critical Rayleigh number decreases. Thus Da has dual effect on oscillatory and stationary convection. The variation of R_{Tc} with Ta for different values of solute Rayleigh number R_s on the onset criteria is shown in Figs. 5d and 6d. We observe from these figures that R_{Tc} increases with the increasing R_s . Thus the effect of R_s is to inhibit the onset of convection for both stationary and oscillatory modes. In Figs. 5e and 6e, the variation of R_{Tc} with Ta for different values of Lewis number Le is shown for the fixed values of other parameters. It is observed that R_{Tc} , increases with the increase of Le , for the stationary convection while decreases for the oscillatory convection, indicating that Le has stabilizing effect on stationary convection while destabilizing effect on oscillatory convection.

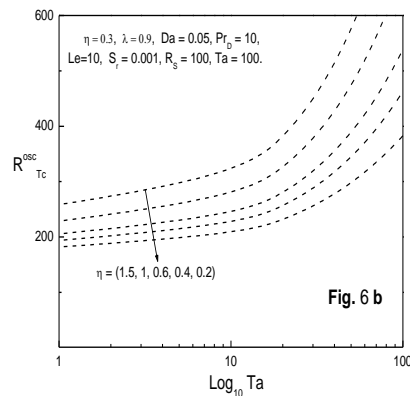


Fig. 6b. Variation of oscillatory critical curves for different values of η

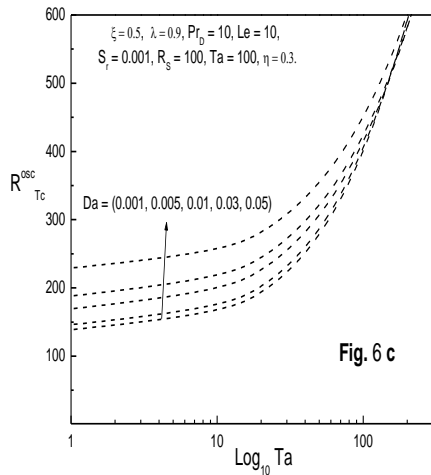


Fig. 6c. Variation of oscillatory critical curves for different values of Da

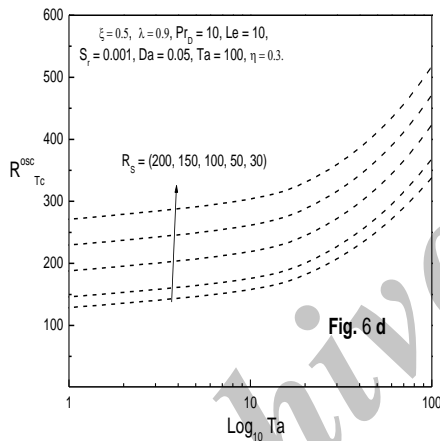


Fig. 6d. Variation of oscillatory critical curves for different values of R_S

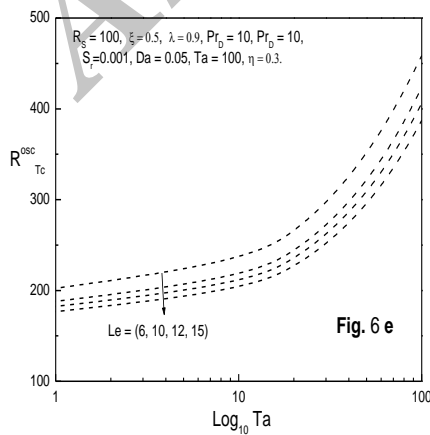


Fig. 6e. Variation of oscillatory critical curves for different values of Le

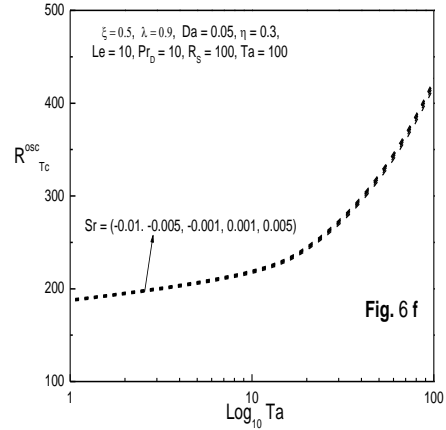


Fig. 6f. Variation of oscillatory critical curves for different values of Sr

The variation of R_{Tc} with Ta for different values of Soret parameter is presented in the Figs. 5f and 6f. We find that as Sr increases positively, the critical values of Rayleigh number and the corresponding wavenumber decrease for stationary mode while as Sr increases negatively, those increase. And the effect of Sr is reversed for oscillatory mode. Thus positive Sr has destabilizing effect for stationary mode and stabilizing effect for oscillatory mode.

5. CONCLUSIONS

The effect of rotation on the onset of double-diffusive convection in a horizontal sparsely packed anisotropic porous layer, in the presence of Soret effect is investigated theoretically using the linear stability theory. The usual normal mode technique is used to solve the linear problem. The applications of the present study has recently increased in view of its potential occurrence in nature and wide range of applications such as high-quality crystal production, liquid gas storage, migration of moisture in fibrous insulation, transport of contaminants in saturated soil, solidification of molten alloys, and geothermally heated lakes and magmas, underground disposal of nuclear wastes, liquid re-injection, electro-chemical and drying processes

Finally, from the results we conclude that the mechanical anisotropy parameter ξ has stabilizing effect on stationary and oscillatory. However the convection sets in as oscillatory mode prior to the stationary mode. The effect of thermal anisotropy parameter η is to inhibit the onset of stationary and oscillatory convection. The Taylor number Ta has a stabilizing effect on the double diffusive convection in sparsely packed anisotropic porous medium. The effect of Darcy number Da is to inhibit the on the onset of stationary convection while it has dual effect on oscillatory convection. The effect of solute Rayleigh number is to delay

both stationary and oscillatory convection. And the effect of Lewis number is to delay the onset of stationary convection while it advances the oscillatory convection. The Darcy Prandtl Pr_D has a dual effect on the oscillatory mode. The Soret parameter has stabilizing effect on oscillatory convection and destabilizing effect on stationary convection.

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