# Realistic Modelling of the Seismic Input: Site Effects and Parametric Studies

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**ABSTRACT:** The work done in the framework of a large international cooperation, showing the very recent numerical experiments carried out within the framework of the EC project "Advanced methods for assessing the seismic vulnerability of existing motorway bridges" (VAB) to assess the importance of non-synchronous seismic excitation of long structures have been illustrated. The definition of the seismic input at the Warth bridge site, i.e. the determination of the seismic ground motion due to an earthquake with a given magnitude and epicentral distance from the site, has been done following a theoretical approach. In order to perform an accurate and realistic estimate of site effects and of differential motion it is necessary to make a parametric study that takes into account the complex combination of the source and propagation parameters, in realistic geological structures. The results for the final local model, characterized by an exaggeratedly thick and low velocity layer, demonstrate that a deep source excites lower frequencies than a shallow one and that the effect of increasing the epicentral distance is to attenuate high frequencies, making the resonant peaks, present at frequencies around 0.8 Hz, the dominant features of the entire spectra. The main practical conclusion of our analysis, verified by laboratory experiments, is that the Warth bridge is likely to well stand the most severe seismic input compatible with the seismic regime of the Eastern Alps.

**Keywords:** Synthetic seismograms; Site effects; Seismic input; Ground motion; Response spectra

# 1. Introduction

It is well accepted that one of the most important factors influencing the space variability of the ground motion is the site response. The local amplification or de-amplification effects can dominate the groundshaking response whenever severe lateral heterogeneities are present in the vicinity of a site. In presence of lateral heterogeneities, like topographic features and/or soft sedimentary basins, the insurgence of local surface waves and local resonances can give rise to a complicated pattern in the spatial ground-shaking scenario, down to a length scale comparable with the smallest wavelength contained in the seismic wave train.

A better understanding of the ground motion spatial

variability can be obtained installing local, dense, seismic arrays at different sites. This implies the recording, with a network of instruments, of multiple seismic sources and the cost of such an operation is evident. The theoretical approach, based on the computation of synthetic seismograms, for the estimation of the site responses uses computer codes, developed from a detailed knowledge of the seismic source process and of the propagation of seismic waves, that can simulate the ground motion associated with the given earthquake scenarios. In such a way, synthetic signals, to be used as seismic input in a subsequent engineering analysis, e.g. for the design of earthquake-resistant structures or for the estimation of differential motion, can be produced at a very low cost/benefit ratio [1]. The realistic modeling of ground motion requires the simultaneous knowledge of the geotechnical, lithological, geophysical parameters and topography of the medium, on one side, and tectonic, historical, paleoseismological, seismotectonic models, on the other, for the best possible definition of the probable seismic source.

We present an example of the theoretical procedure applied to the seismic hazard assessment of the Warth bridge, near Vienna (Austria), where no seismic records are available. The initial stage of the work was thus devoted to the collection of all available data concerning the deep and shallow geology, the construction of cross-sections along which to model the ground motion, and the specification of the possible seismic sources. Following the upgrade and the improvement of the initial databank of seismic sources (i.e. focal mechanisms) and structural (i.e. bedrock and local) models, the seismic input calculation has been performed at different stages of parametric studies, adopting a set of possible scenarios for the seismic source-Warth bridge configurations.

### 2. First Parametric Study

# 2.1. Definition of Bedrock and Local Struc u u Models

The regional structural model for the rea bere the

Warth site lies is adapted from the I-dataset [2]. The vertical dependence of the elastic and anelastic parameters is shown in Figure (1). Starting from the available Warth bridge section plan, see Figure (2a), a digitized model of the geological cross-section underlying the bridge has been assembled, see Figure (2b). On the basis of the geological and geotechnical information available and considering the results obtained from a local refraction seismic survey, the elastic and the anelastic parameters, see Table (1), have been assigned to the various polygons, corresponding to the different geotechnical units, contained in the section, see Figure (2b).



Figure 1. Vertical dependence of the elastic (density, P and S wave velocity) and anelastic ( $Q_p$  and  $Q_s$ ) parameters of the average regional model assumed for the Warth area.



Figure 2. a) Warth bridge section plan. b) Local heterogeneous model along Warth bridge and its geotechnical units. Black triangles show the sites of the abutments and of the piers and their relative distance along the section.

Unit	Density g/cm <sup>3</sup>	P-Wave Velocity km/s	$Q_P$	S-Wave Velocity km/s	Qs
1	1.5	0.30	40.0	0.20	15.0
2	1.7	0.49	40.0	0.25	15.0
3	2.0	0.70	50.0	0.26	20.0
4	1.8	0.70	50.0	0.29	20.0
5	2.3	0.80	50.0	0.30	20.0
6	2.3	0.80	50.0	0.40	20.0
7	1.8	1.70	50.0	0.50	20.0
8	2.3	2.10	150.0	1.00	60.0
9	2.3	3.00	150.0	1.90	60.0
10	2.2	1.80	100.0	1.10	40.0

 
 Table 1. Elastic and anelastic parameters of the geotechnical units shown in Figure (2b).

#### 2.2. Definition of Source Models

To define the possible seismic sources that control the seismic hazard of the Warth region, we used the available database of focal mechanisms [3]. Taking into account the magnitudes and the epicentral distances from the Warth region, we initially selected the five sources, whose focal mechanisms are shown in Figure (3) and their parameters listed in Table (2).

The distances of the selected sources from the Warth bridge site (assumed geographical coordinates Latitude =  $47.660^{\circ}N$  and Longitude =  $16.170^{\circ}E$  are respectively 41.2km, 20.3km, 26.8km, 8.6km, and 13.7km. As a conservative choice, mag itude



Figure 3. Focal mechanisms of the 5 events reported in Table (2) and Warth site (triangle).

Table 2. Focal mechanisms for the five selected sources.

Source	Lon E	Lat N	Focal	Strike	Dip	Rake	Magnitude
id	(°)	(°)	Depth	(°)	(°)	(°)	Ms (Mb)
			(km)				
SEM63	16.200	48.030	?	180	20	90	?
SEM64_1	15.920	47.730	3	90	81	311	(4.7)
SEM64_2	15.950	47.850	1	100	70	31	(5.4)
SEE72	16.120	47.730	18	190	70	324	5.5 (4.9)
NEU72	16.020	47.730	19	127	80	190	4.4

(equal to 5.5) and hypocentral depth (equal to 5km) have been kept constant for all the sources, and the source finiteness has been taken into account by properly weighting the source spectrum using the scaling laws of [4], as reported in [5]. The synthetic seismograms at the base of each pier (displacements, velocities and accelerations for the radial, transverse and vertical components) have been computed, with cut-off frequency at 10Hz, using the bedrock model. From the analysis in time (amplitude and duration) and frequency domain, we obtain that source *SEE*72, see Table (2), is the most interesting from the seismic hazard assessment point of view. Therefore *SEE*72 has been used for the preliminary computation of the seismic input at the the other the bridge site.

A paramet. stud, of the ground motion was performed ford, to take into account the variations due to the chick of the seismic source, different groun for the seismic source, different groun for the warth site have been studied, in order to reach the maximum excitation in both fongitudin, rand transverse direction. Starting from the N ximum Historical Earthquake, an additional study has been made considering the magnitude correspond-1 g to both the Maximum Credible Earthquake and the Maximum Design Earthquake.

Starting from the source model *SEE*72, in the first three tests the source depth is fixed at 5km (as a conservative choice) while the distance to the bridge is 8.6km. The strike has been varied from 0° to 360°, see Figure (4), the dip from 0° to 90°, see Figure (5), and the rake from 0° to 180°, see Figure (6). For each test, displacement, velocity and acceleration are computed assuming the other angles fixed at the values shown in Table (2).

For a more general study, three other tests have been performed: the dip angle has been fixed at  $45^{\circ}$ ,  $70^{\circ}$  and  $90^{\circ}$  varying the strike and the rake, see Figure (7). Once established the most effective combination for the transverse and the radial components of motion (strike =  $60^\circ$ , dip =  $70^\circ$ , rake = 0,  $90^\circ$ for transverse and radial components respectively), in the last test the source depth (from 1 to 20km) and the epicentral distance (from 5 to 20km) have been varied, see Figure (8). The results shown in Figure (8) allow us to conclude that, for a fixed magnitude, the most effective focal depth, for an epicentral distance of 8-9km, is 6km both for the transverse, see Figure (8a), and the radial, see Figure (8b) components of motion (the focal depths between 1 and 3km are shown for completeness but they are

F. Romanelli, et al



**Figure 4.** Polar plot of the maximum amplitude of the ground motion - a) acceleration (*cm*/*s*<sup>2</sup>); b) velocity (*cm*/*s*); c) displacement (*cm*) - versus the strike angle, for the three components - transverse (squares); radial (circles); vertical (triangles). The rake (324°) and dip (70°) angle values are those of mechanism *SEE*72 reported in Table (2). The thick line represents the case with strike angle (190°).



**Figure 5.** Polar plo f the maximum amplitude of the ground nouon - a cceleration (cm/s<sup>2</sup>); b) velocity (cm/s); c) displacen at (cm) - versus the rake angle, for the mponents - transverse (squares); radial (circles); ertical (triangles). The strike (190°) and dip (70°) angle values are those of mechanism SEE72 reported in Table (2). The thick line represents the case with rake angle (324°).



**Figure 6.** Polar plot of the maximum implitude of the ground motion-a) acceleration (cm/s<sup>2</sup>); b) velocity (cm/s); c) displacement (cm)-versus the dip angle 1...th, bry e components-transverse (squares); radial (circles); vertical (triangles). The strike (190°) and rake (324°) ang. values are those of mechanism SEE72 reported in Table (2). The thick line represents the case with dip angle (70°).



Figure 7. Plot of the maximum amplitude of the ground motion (acceleration) versus the strike and rake angle, for the three components: a) transverse; b) radial; c) vertical. The dip angle is 70°.



**Figure 8.** Plot of the maximum amplitude of the ground motion (acceleration -cm/s<sup>2</sup>) -versus epicentral distance and source depth, for the three components: a) transverse-rake angle = 0°; b) radial-rake angle = 90°; c) vertical-rake angle = 90°, strike = 60° and dip =70°.

highly unrealistic).

The most effective focal mechanisms in radiating *SH* waves (transverse component) and *P-SV* waves (radial and vertical component), propagating towards the Warth bridge in the adopted structural model are listed in Table (3).

Table 3. Focal mechanisms for the selected sources.

Source	Focal Depth	Strike	Dip	Rake	Magnitude
id.	(km)	(°)	(°)	(°)	
SEE72	5	190	70	324	5.5
SEE-SH	5 and 6	60	70	0 and180	5.0; 5.5; 6.0
SEE-PSV	5 and 6	60	70	90 and 270	5.0 5. ; 6.0

#### 2.3. Definition of the Seismic I. ut: "sults from the First Parametric Study

To deal both with realistic so, e a structural models, including topographi, f, 'ures, a hybrid method has been develope. the combines modal summation and the finite di ferei ce technique (e.g. [6, 7]), and optimizes the se on the advantages of both methods. Way propag, ion is treated by means of the modal summa. technique from the source to the vicinity of Vie local, heterogeneous structure that we may wat to model in detail. A laterally homogeneous anelastic structural model is adopted, that represents the average crustal properties of the region. The generated wavefield is then introduced in the grid that defines the heterogeneous area and it is propagated according to the finite differences scheme, see Figure (9). With this approach, source, path and site effects are all taken into account, therefore a detailed study of the wavefield that propagates even at large distances from the epicenter, without having to resort to convolutive methods, that may be quite misleading (e.g. [1, 8, 9]), is possible.

In the hyb d sch, he the local heterogeneous model is couple, with the average regional model used in an initial analysis, see Figure (1). The minimum S-v, we velocity in the model is 220m/s, and he mesh sed for the finite differences is defined with a grid spacing of 3m. This allows us to carry out the computations at frequencies as high about 8Hz (since 10 grid points per minimum wa blength are needed), well above the frequency unge relevant for large dimensions objects, like Warm bridge.

The synthetic time signals (displacements, velocities and accelerations) have been calculated for the three components of motion. The working magnitude is 5.5 (seismic moment,  $M_0$ , equal to 1.8  $10^{17}Nm$ ), corresponding to the nearest largest recorded event, but the magnitude range 5.0-6.0 has



Figure 9. Scheme of the hybrid (modal summation plus finite differences scheme) method (from [1]).

been explored. The study of possible directivity effects in the direction of the Warth cross section has been performed adopting a new method, based on the modeling of a Haskell-type [10, 11] source. A stochastic component allows us to build a spectrum (amplitude and phase) of the source function that takes into account both the rupture process and directivity effects. As an example the transverse acceleration time series, calculated at the bridge piers, with the two methods (scaled point source [4] and Haskell-type source), are plotted in Figure (10a) while in Figure (10b) the corresponding Fourier amplitude spectra are shown. Figure (11) shows the same results for the radial component of motion.

We give an estimate of the local response at each site, evaluating the Spectral Ratios, using both the Fourier Spectra (FSR) and the Response Spectra (RSR), corresponding to the laterally varying model and to the reference bedrock model. As an example, the response spectra for the signals (Haskell-type source) of Figures (10a) and (11a), and the corresponding RSR and FSR, are shown in



Figure 10. a) Acceleration time series, for transverse component and source model SEE-SH, calculated at the position of the bridge piers (plus one at the abutment), considering the Gusev scaling law for magnitude 5.5 with (dashed lines) and without (solid lines) directivity effects. Fourier amplitude spectra of the signals shown in a), considering the Gusev scaling law for magnitude 5.5 with (b1) and without (b2) directivity effects, together with those obtained for the bedrock model (b3). The legend is the same for the three sets of Fourier spectra.

Figures (12) and (13) respectively. In Figure (14), the *RSR* versus epicentral distance and frequency are shown for transverse and radial components of motion, together with the results of the radial/vertical (henceforth named H/V) *RSR*.

# 3. Second Parametric Study

### 3.1. Definition of Source Models

In the frequency domain, the accelerograms shown in Figures (10) and (11) exhibit the greatest amplitudes in the frequency range from 3 to 6Hz, reaching considerable peak values (around  $400cm/s^2$ ). Another parametric study has been performed in order to find a seismic source-Warth site configuration providing a set of signals whose frequency content is concentrated around 1Hz. Actually, 1Hz is the frequency that corresponds approximately to that of the fundamental transverse mode of oscillation of the bridge and we focussed our analysis on the transverse component of motion.

The computation of synthetic seismograms



Figure 11. a) Acceleration time series, for transverse component and source model SEE-SH, calculated at the position of the bridge piers (plus one at the abutment), considering the Gusev scaling law for magnitude 5.5 with (dashed lines) and without (solid lines) directivity effects. Fourier amplitude spectra of the signals shown in a), considering the Gusev scaling law for magnitude 5.5 with (b1) and without (b2) directivity effects, together with those obtained for the bedrock model (b3). The legend is the same for the three sets of Fourier spectra.

F. Romanelli, et al



Figure 12. a) Response spectra for the signals of Figure (10a) (with tirec. ity) and b) the corresponding RSR and c) FSR. The legend is the same.



Figure 13. a) Response spectra for the signals of Figure (11a) (with directivity) and b) the corresponding RSR and c) FSR. The legend is the same.



Figure 14. RSR for a) the transverse component, b) radial component and c) H/V, versus epicentral distance and frequency.

(accelerations for the transverse component) has been carried out considering the source *SEE-SH* t ir ed in the bedrock model. The focal depth and the ep ep ral distance have been varied in the ing. 5-20km and 5-100km, respectively. The result of Figure (15) show that a relevant value c PC (e.g greater than  $100cm/s^2$ ) in the period range of 1. crest (0.8-1.2s) can be reached by a, geophysically bund, source 12kmdeep at an epicentral distance of 30km.

#### 3.2. Definition of the S. Smic Input

The same computation is and analysis described in Section 2.3, but 1 mited to the transverse component of motion, have been carried out for the configuration defined in Section 3.1 (i.e. strike =  $60^{\circ}$ ; dip =  $70^{\circ}$ ; rake =  $0,180^{\circ}$ ; depth = 12km; epicentral distance = 30km). In Figure (16) the Fourier amplitude spectra of the acceleration time series calculated for a magnitude equal to 6.0, with and without directivity effects, are shown. The results show that, even if the seismic energy around 1Hz can be relevant (see bedrock curves), the local structure beneath the Warth bridge greatly amplifies the frequency components between 3 and 7Hz, i.e. a frequency range not corresponding to the fundamental transverse







Figure 16. Fourier amplitude spectra of the transverse accelerations calculated at the 8 sites shown in Figure (2a), for a focal depth of 12km, a source distance of 30km, a magnitude equal to 6.0, with (a) and without (b) directivity effects. The thick black lines correspond to the curve for the bedrock model (common for all the eight sites).

mode of oscillation of the bridge (about 0.8Hz). In Figure (17) the *RSR* versus epicentral distance and frequency is shown.



Figure 17. RSR versus epicentral distance and frequency.

#### 4. Final Definition of the Seismic Input

In the final set of computations, the local heterogeneous model has been iteratively changed. Initially, the S-wave velocities of the uppermost units (units 1-7 of Figure (2b)) have been halved (see for comparison Figure (18a) and Figure (2b)); then, in order to characterize the local structure with lower resonant frequencies, some of the geotechnical v. its have been assigned to a class characterized by 1 wer velocities (unit 7 to unit 4 in Figure (18b); um 7 unit 8 to unit 4 in Figure (18c)).

The synthetic time series have been c mputed for the transverse component of  $n_{1}$  ion sing he focal mechanism obtained from the para etc. studies and



Figure 18. Local geotechnical models of Warth bridge section obtained lowering successively the S-wave velocities of the uppermost units.

the configurations described in Sections 3 and 5: SS1) Strike =  $60^\circ$ ; Dip =  $70^\circ$ ; Rake =  $0^\circ$ ; Depth =

- 5km; Distance = 8km, Magnitude = 5.5
- SS2) Strike =  $60^\circ$ ; Dip =  $70^\circ$ ; Rake =  $0^\circ$ ; Depth = 12km; Distance = 30km, Magnitude = 6.5

In such a way, using the local models, a, b and c, shown in Figure (18), we obtain six different source-local structure pairs, henceforth named: SS1a; SS1b; SS1c; SS2a; SS2b; SS2c. In Figure (19) the Fourier amplitude spectra are shown, for the acceleration time series calculated for SS1a, SS1b, SS2b, SS1c and SS2c configurations. In the model of Figure (18a), only the first two sites have significant resonance at frequencies lower than 3Hz. Using the model shown in Figure (18b), at the s. from 3 to 6, due to the increased thicknes. If the ver with S-wave velocity equal to 150m; rest ance at lower frequencies, i.e. between 1 and  $\forall z$ , is isible. As expected, Figures (19b) an 1(19c) show that a deep source excites lower e than a shallow one, and that the frequent effect of increasing the epicentral distance is to attenuate vigh frequencies. Similar conclusions may be drawn from *igures (19d)* and (19e), with the difference that, in his case, the frequency peaks appear at significantly bw. frequencies. In particular, site 3 is characterized by a peak spectral acceleration (about  $400 cm/s^2$ ) very near to the target frequency (i.e. 0.8Hz). The acceleration time series are shown in Figure (20). In Figure (21) the RSR versus epicentral distance and frequency are shown for configurations SS1a, SS1b, SS2b, SS1c and SS2c. The nonlinear effects are not treated in the following; we just mention that the assumption of linearity between stress and strain can be no longer valid for accelerations larger than  $200cm/s^2$  (e.g. [12]). Due to nonlinearity the actual shear wave velocity decreases with increasing stress, and hysteresis leads to energy loss at any deformation cycle. As a consequence, the resonance of surficial layers can be shifted to lower frequencies, and this can lead to a lower amplification of ground motion at higher frequencies.

#### 5. Discussion and Conclusions

Two parametric studies of the ground motion have been performed, taking into account the variations due to the choice of the focal mechanism parameters and the geometry of the seismic source. Different ground motions at the Warth site, which are consistent both with the Maximum Credible Earthquake and with the Maximum Design Earthquake, have been studied in order to define the maximum excitation in the



**Figure 19.** Fourier amplitude spectra of the cansverse accelerations, calculated at the eight pier sites, using the source-section configuration a) SS1a ) S b; c, 2b; d) SS1c; and e) SS2c.



Figure 20. Transverse acceleration time series corresponding to configuration SS2c, calculated at the eight pier sites. The amplitude of the signals is normalized with respect to the maximum one.



Figure 21. RSR versus epi, htral distance and frequency for configuration: a) SS1a; b) SS1b; c) SS2b; d) SS1c; e) SS2c.

transverse direction  $C_i$  the oridge. With the parametric study we have define i a seismic source-Warth site configuration that provides a set of signals whose seismic energy is concentrated around 1Hz, frequency that corresponds approximately to that of the fundamental transverse mode of oscillation of the bridge. The results have led to the definition of two possible scenarios:

- S1) Strike =  $60^{\circ}$ ; Dip =  $70^{\circ}$ ; Rake =  $0,180^{\circ}$ ; Depth = 5km; Distance = 8km, Magnitude = 5.5.
- S2) Strike =  $60^{\circ}$ ; Dip =  $70^{\circ}$ ; Rake =  $0,180^{\circ}$ ; Depth = 12km;Distance = 30km, Magnitude = 6.5.

The analysis of the computed seismic input has been carried out in the time domain (broad band ground motion time series) and in the frequency domain (Fourier and Response Spectra). We give an estimate of the local response at each site, evaluating the Spectral Ratios, using both the Fourier Spectra (*FSR*) and Response Spectra (*RSR*) for the laterally varying model normalized to the ones computed for the bedrock model.

In the final set of computations, the local heterogeneous model has been iteratively changed, in order to characterize the local structure with lower resonant frequencies: a) the S-wave velocities of the original uppermost units have been halved; then, b) some of the geotechnical units have been assigned to a class characterized by lower velocities. The results for the final local model, characterized by an exaggeratedly thick and low velocity layer, demonstrate that a deep source excites lower frequencies than a shallow one and that the effect of increasing the epicentral distance is to attenuate high frequencies, making the resonant peaks, present at frequencies around 0.8*Hz*, the dominant features of the entire spectra. Therefore the main practical conclusion of our analysis, verified by laboratory experiments carried out at *JRC-ISPRA* within the *VAB* Project [13], is that the Warth bridge is likely to well stand the most severe seismic input compatible with the seismic regime of the Eastern Alps.

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The *GMT* software [14] has been used for the preparation of Figures (1) and (2).

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